Simulating rotating convection at very small but finite Ekman numbers



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Overview

- 1) Introduction
- 2) Non-hydrostatic quasi-geostrophic equations (NHQG)
- 3) Rescaled incompressible Navier-Stokes equations (RiNSE)
- 4) Numerical results
- 5) Conclusions and Outlook

1) Introduction

Convection drives fluid flows on Earth & beyond

Planetary interiors



Oceans on Earth ...

Planetary atmospheres







... and beyond





Paradigm: Rotating Rayleigh-Bénard convection



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The astrophysical limit: $Ek \rightarrow 0$, $Re_H \rightarrow \infty$



Rotating faster State-of-the-art numerical and experimental techniques cannot reach realistic Ek

$\underline{\text{Lowest } Ek \text{ achieved:}}$

DNS (plane layer geometry): $Ek \geq 5 \times 10^{-9}$ [Song, Shishkina, Zhu, JFM 2024 a, b]

DNS (spherical geometry): $Ek \geq 10^{-7} \label{eq:Ek}$ [Gastine and Aurnou, JFM 2023]

Laboratory experiments (Eindhoven): $Ek \ge 5 \times 10^{-9}$ [Cheng+al. *GAFD* 2018, Kunnen *JT* 2020]

Key Challenges

- Fast inertial waves with $O(Ek^{-1})$ frequencies \rightarrow prohibitive time stepping requirements
- Columnar structures with aspect ratio $\sim Ek^{-1/3} \gg 1$ (Taylor-Proudman) + Ekman layers of depth $O(Ek^{1/2})$ are difficult to resolve
- Scale disparity causes ill-conditioned matrices in discretized problem
 → numerical errors

How to tackle these?

Linear stability theory:

[Chandrasekhar (1953)]

Linearize about conduction state $\mathbf{u} = \mathbf{0}, \theta = 0$ to obtain the system $\partial_t \mathbf{u} + \frac{1}{Ek} \hat{\mathbf{z}} \times \mathbf{u} = -\nabla p + \frac{Ra}{Pr} \theta \hat{\mathbf{z}} + \nabla^2 \mathbf{u}$ $\partial_t \theta = \hat{\mathbf{z}} \cdot \mathbf{u} + \frac{1}{Pr} \nabla^2 \theta$ $\nabla \cdot \mathbf{u} = 0$ which admits solutions of the form $(\mathbf{u}, \theta) = (\mathbf{U}(z), \Theta(z)) \exp(ik_x x + ik_y y + \sigma t)$

which admits solutions of the form $(\mathbf{u}, \mathbf{v}) = (\mathbf{v}, \mathbf{v}, \mathbf{v})$ For Pr > 0.68: **Onset:** $Ra_c^{Ek \to 0} 3\left(\frac{\pi^2}{2}\right)^{2/3} Ek^{-4/3} \approx 8.7 Ek^{-4/3}$ $\rightarrow \text{Rotation suppresses convection}$

> Critical horiz. wavenumber: $k_c^{Ek \rightarrow 0} (2\pi^4)^{1/3} Ek^{-1/3} H^{-1}$ \rightarrow Thin columns

2) Nonhydrostatic quasi-geostrophic equations

Sprague et al., J. Fluid Mech. 2006

Change of units

• Let $\epsilon \equiv Ek^{1/3}$ and define $\widetilde{Ra} = \epsilon^4 Ra$ (cf. linear stability analysis).

• Measure lengths in units of $\ell = \epsilon H$, time in units of $\tau_{\nu} = \ell^2 / \nu = \epsilon^2 H^2 / \nu$, with velocity scale $U = \frac{\ell}{\tau_{\nu}} = \epsilon^{-1} \nu / H$, such that

$$\nabla \to \frac{1}{\epsilon} \nabla, \qquad D_{t} \to \frac{1}{\epsilon^{2}} D_{t}, \qquad \boldsymbol{u} \to \frac{1}{\epsilon} \boldsymbol{u}, \qquad p \to \frac{1}{\epsilon^{3}} p$$

leading to

$$D_{t}\mathbf{u} + \frac{1}{\epsilon}\mathbf{\hat{z}} \times \mathbf{u} = -\frac{1}{\epsilon}\nabla p + \frac{1}{\epsilon}\frac{\widetilde{Ra}}{Pr}\theta\mathbf{\hat{z}} + \nabla^{2}\mathbf{u}$$
$$D_{t}\theta = w + \frac{1}{Pr}\nabla^{2}\theta$$
$$\nabla \cdot \mathbf{u} = 0$$

Multi-scale asymptotic expansion

Consider $\epsilon = Ek^{1/3} \rightarrow 0$, assume $\widetilde{Ra} = RaEk^{4/3} = O(1)$

- 1) Introduce slow variables $Z = \epsilon z, T = \epsilon^2 t$ in addition to fast variables x, y, z, tsubstituting $\partial_z \to \partial_z + \epsilon \partial_Z, \quad \partial_t \to \partial_t + \epsilon^2 \partial_T$
- 2) Decompose $\boldsymbol{v} = (\boldsymbol{u}, p, \theta)$ into $\boldsymbol{v} = \overline{\boldsymbol{v}}(Z, T) + \boldsymbol{v}'(\boldsymbol{x}, Z, t, T)$, with

$$\overline{f}(Z,T) \equiv \lim_{L,\tau\to\infty} \frac{1}{\tau \operatorname{Vol}(D)} \int_D \int_0^\tau f(\mathbf{x}, Z, t, T) d\mathbf{x} dt.$$

3) Expand $w = (\overline{v}, v')$ as $w = w_0 + \epsilon w_1 + \epsilon^2 w_2 + O(\epsilon^3)$, solve order by order

- 4) Leading-order geostrophic balance: $\hat{\mathbf{z}} \times \mathbf{u}'_0 = -\nabla p'_0$ Taylor-Proudman constraint: $\partial_z \mathbf{u}'_0 = \partial_z p'_0 = \partial_z \theta'_1 = 0$
- 5) At next order: avoid secular growth of w_2 with z by solvability condition
 - \rightarrow closed set of equations

Non-hydrostatic quasi-geostrophic equations

$$D_t^{\perp} w + \frac{\partial}{\partial Z} \left((\nabla_{\perp}^2)^{-1} \omega \right) = \frac{\widetilde{Ra}}{Pr} \theta' + \nabla_{\perp}^2 w \qquad (1)$$

$$D_t^{\perp}\omega - \frac{\partial w}{\partial Z} = \nabla_{\perp}^2\omega \tag{2}$$

$$D_t^{\perp}\theta' + w\partial_Z \,\overline{\theta} = \frac{1}{Pr} \nabla_{\perp}^2 \theta' \tag{3}$$

$$\partial_T \overline{\theta} + \partial_Z \overline{w\theta'} = \frac{1}{Pr} \partial_Z^2 \overline{\theta} \tag{4}$$

$$\partial_x u + \partial_y v = 0. \tag{5}$$

Where $D_t^{\perp} = \partial_t + u\partial_x + v\partial_y$, $\nabla_{\perp}^2 = \partial_x^2 + \partial_y^2$, and the vorticity $\omega = \partial_x v - \partial_y u$.

Key features:

Small parameter ϵ & fast scale z are eliminated.

→ Filtering fast waves with $O(Ek^{-1})$ frequencies and Ekman layers Eqs. (1-5), a.k.a. **3D Hasegawa-Mima Eqs**., related equations were studied by **Edriss Titi** + al.

- J. Evol. Equ. 21 (2021), 2923–2954.
- J. Diff. Equ. 269.10 (2020): 8736-8769.
- J. Math. Phys. 59.7 (2018)
- Comm. Math. Phys. 319 (2013): 195-229.

Global existence & uniqueness, cont. dependence on initial data of solutions to (1-5) with small vertical dissipation in (2).

Equivariance under reflections $\begin{cases} (x, y) \rightarrow (-x, y) \\ (u, v, w, \omega) \rightarrow (-u, v, -w, -\omega) \\ (\theta', \overline{\theta}) \rightarrow (-\theta', \overline{\theta}) \\ \text{and similar for } (x, y) \rightarrow (x, -y) \text{ indicates absence of handedness.} \end{cases}$

3) Rescaled incompressible Navier-Stokes equations

Julien, **AvK** et al., arXiv:2410.02702 (2025) **AvK**, Julien et al., arXiv:2409.08536 (2025)

Rescaled Navier-Stokes equations (RiNSE)

$$D_{t}\mathbf{u} + \frac{1}{\epsilon}\mathbf{\hat{z}} \times \mathbf{u} = -\frac{1}{\epsilon}\nabla p + \frac{1}{\epsilon}\frac{\widetilde{Ra}}{Pr}\theta\mathbf{\hat{z}} + \nabla^{2}\mathbf{u}$$
$$D_{t}\theta = w + \frac{1}{Pr}\nabla^{2}\theta$$
$$\nabla \cdot \mathbf{u} = 0$$

Starting point:

Boussinesq-Navier-Stokes equations ($D_t \equiv \partial_t + \boldsymbol{u} \cdot \boldsymbol{\nabla}$)

Idea: rescale full Boussinesq-Navier-Stokes equations based on NHQG scaling laws valid at $Ek \rightarrow 0$

Two key steps:

1) Adopt anisotropic length scales, such that

$$\nabla^2 \to \nabla^2_{\epsilon} \equiv \partial_x^2 + \partial_y^2 + \epsilon^2 \partial_z^2$$

and material derivatives become

$$D_t \to D_t^{\epsilon} \equiv \partial_t + u \partial_x + v \partial_y + \epsilon w \partial_z = \partial_t + \boldsymbol{u}_{\perp} \cdot \boldsymbol{\nabla}_{\perp} + \epsilon w \partial_z$$

2) Decompose temperature as $\theta = \Theta(z, t) + \epsilon \vartheta(x, y, z, t)$.



Grooms et al. PRL (2010) 16

Rescaled Navier-Stokes equations (RiNSE)

Equations governing horizontal fluctuations $(k_x, k_y) \neq (0,0)$

$$D_t^{\epsilon} \mathbf{u} + \epsilon^{-1} \mathbf{\hat{z}} \times \mathbf{u} = -\epsilon^{-1} \nabla_{\epsilon} p + \frac{Ra}{Pr} \vartheta \mathbf{\hat{z}} + \nabla_{\epsilon}^2 \mathbf{u}$$
$$D_t^{\epsilon} \vartheta + w(\partial_z \Theta - 1) = \frac{1}{Pr} \nabla_{\epsilon}^2 \vartheta$$

Where $\epsilon = Ek^{1/3}$ and the *reduced* Rayleigh number is $\widetilde{Ra} = Ra Ek^{4/3} = Ra \epsilon^4$.

Horizontally averaged temperature is governed by

$$\epsilon^{-2}\partial_t\Theta + \partial_z\overline{w\vartheta}^{xy} = \frac{1}{Pr}\partial_{zz}\Theta$$

Becomes negligible as horizontal domain size is increased

4) Numerical results

Julien, **AvK** et al., arXiv:2410.02702 (2025) **AvK**, Julien et al., arXiv:2409.08536 (2025)

Numerical method

- Use spectral PDE solver Coral [Miquel JOSS 2021]
- Expand fields in Fourier basis in x, y and in Chebyshev polynomials in z

$$(\mathcal{M}\partial_t - \mathcal{L}_I)\mathbf{v} = \mathcal{N}(\mathbf{v}, \mathbf{v})$$

 $\mathbf{v} = \mathbf{h}^{(\mathbf{v}, \mathbf{v})}$
 $(\mathbf{M}\partial_t - \mathbf{L}_I)\mathbf{v} = \mathbf{b}$

- Assess **numerical accuracy** at low *Ek* in two complementary ways:
 - 1. Compute condition number of matrices on LHS
 - 2. Consider generalized eigenvalue problem $(\mathbf{M}s \mathbf{L}_I)\mathbf{v} = 0$

Conditioning properties at low *Ek*

Condition number of discretized linear operator:

 $\kappa(\mathbf{L}_{I}) = \|\mathbf{L}_{I}\|_{2} \|\mathbf{L}_{I}^{-1}\|_{2}$

Large κ indicates sensitivity to errors in right-hand side



RiNSE formulation ensures adequate preconditioning





Nonlinear, direct numerical simulations

Integrate the RiNSE with stress-free BC from small-amplitude noise

• over a wide range of Ekman numbers

 $10^{-1} \le Ek \le 10^{-15}$

• and

Pr = 1

- in a slender domain of horizontal extent $L = 10\ell_c \approx 48.2 \text{Ek}^{1/3}H$
- Resolve thermal boundary layers in vertical, energy dissipation scale in horizontal

Spoiler:

Will present evidence for convergence of full governing equations towards the reduced NHQGE at small *Ek*.

Validation of RiNSE at moderate *Ek*



RiNSE reproduces Nusselt numbers from the literature

$\overline{\omega_z}(x,y)$ $E_{L_y} = 10^{-15} (\text{Ro} \approx 0.00008) \text{ Ek} = 10^{-8} (\text{Ro} \approx 0.07) \text{ Ek} = 10^{-7} (\text{Ro} \approx 0.15) \text{ Ek} = 10^{-6} (\text{Ro} \approx 0.4)$ -100-50 $L_y/2$ 0 -50-100 $L_x/2$ $L_x/2$ $L_x/2$ $L_x/2$ $L_r 0$ $L_r 0$ 0 2 Vertically averaged vertical vorticity: 1 $\overline{\omega_z}(x,y) \equiv \int_0^1 (\partial_x v - \partial_y u) dz$ $\operatorname{Skew}(\overline{\omega_z})$ 0 $^{-1}$ Data shown for $\widetilde{Ra} = 120$. NHQG -2 - 2 10^{-12} 10^{-15} 10^{-9} 10^{-6} Symmetries of NHQG are respected at $Ek~\lesssim~10^{-9}$ 25 Ek

Cyclone anti-cyclone symmetry breaking







Boundary layer flow morphology ($z < \delta_{\theta}$)



Boundary layer flow morphology ($z < \delta_{\theta}$)



Global statistics in agreement within one standard deviation

Alternative parameter space cut: Ra = const.



[Julien et al. PRL 2012]

Thermal boundary layer depth



Denote by $\delta_artheta$ the location of the local maximum of $artheta_{rms}(z)$

Boundary layer depth converges to NHQG value





Summary

- Described NHQG equations arising from asymptotic expansion
- Introduced RiNSE formulation based on NHQG scaling laws
- Performed numerical simulations at Ek down to 10^{-15}

(six orders of magnitude smaller than previous state of the art)

- Numerical evidence points to convergence of RiNSE to NHQGE at small *Ek* based on symmetry, global statistics, small-scale structure
- Threshold value below which convergence is observed Ek $\approx \widetilde{\text{Ra}}^{-\frac{15}{4}}$ confirming theoretical predictions by [Julien+al. GAFD 2012]



Under review: Julien et al. arXiv:2410.02702, van Kan et al. arXiv:2409.08536



Outlook

- RiNSE opens the door to exploration of unprecedented parameter regimes
 Important additional physics should be included, such as
 - Prandtl numbers different from unity
 - Alternative boundary conditions (e.g., no-slip walls \rightarrow Ekman pumping)
 - Internal heating
 - Latitudinal variations
 - Magnetic fields
 - ...



Thank you for your attention!







