

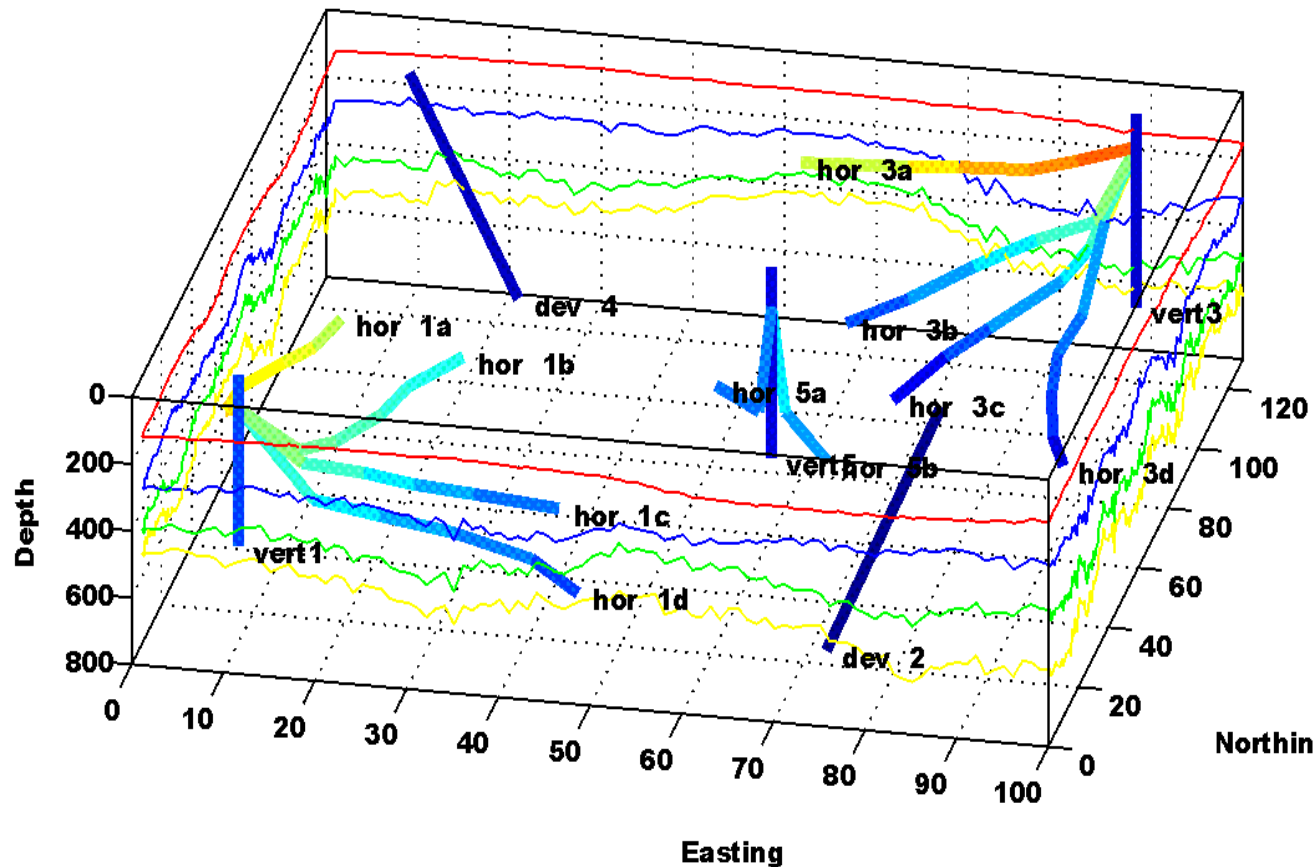
Computational Challenges in Reservoir Modeling

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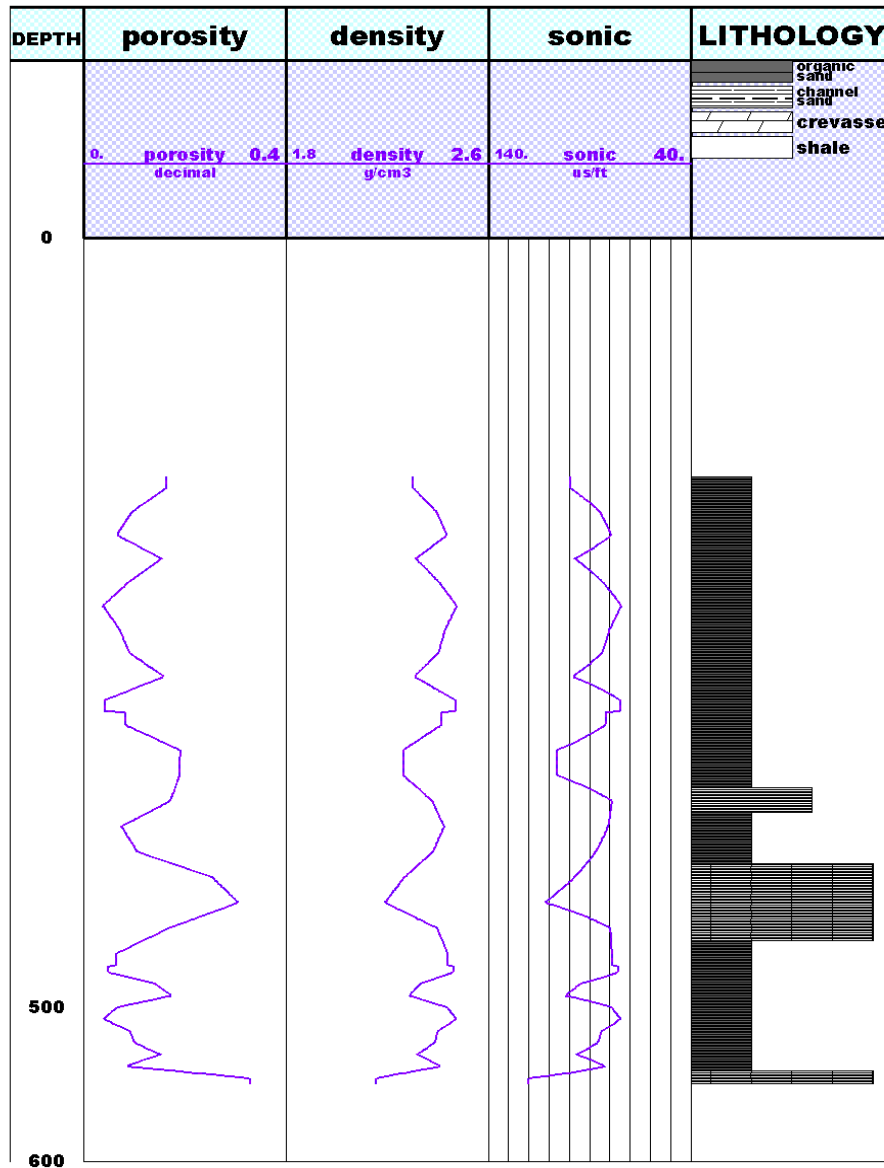
Well Data

3D view of well paths



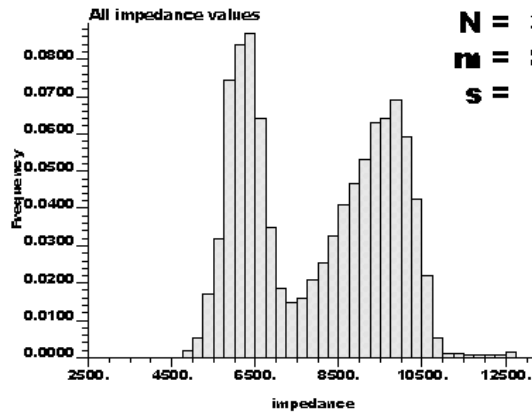
- Inspired by an offshore development
- 4 platforms
- 2 vertical wells
- 2 deviated wells
- 8 horizontal wells

Sample well log

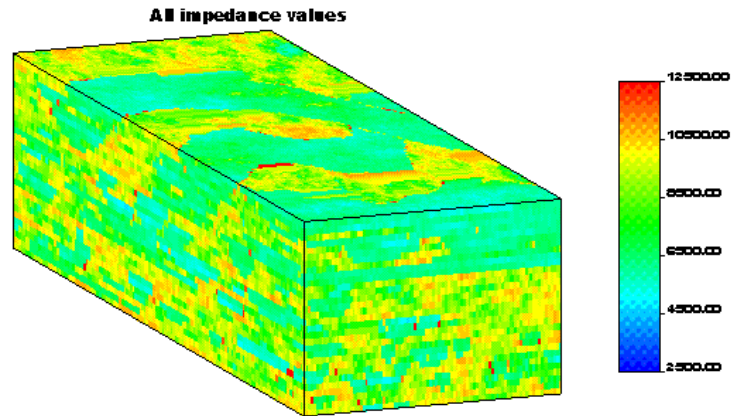


- **Well data include:**
 - horizontal location
 - true vertical depth
 - porosity
 - permeability
 - density
 - sonic log
 - lithology
 - layer indicator
- **Well markers:**
 - true vertical depth when well intercepts a layer surface

Seismic impedance



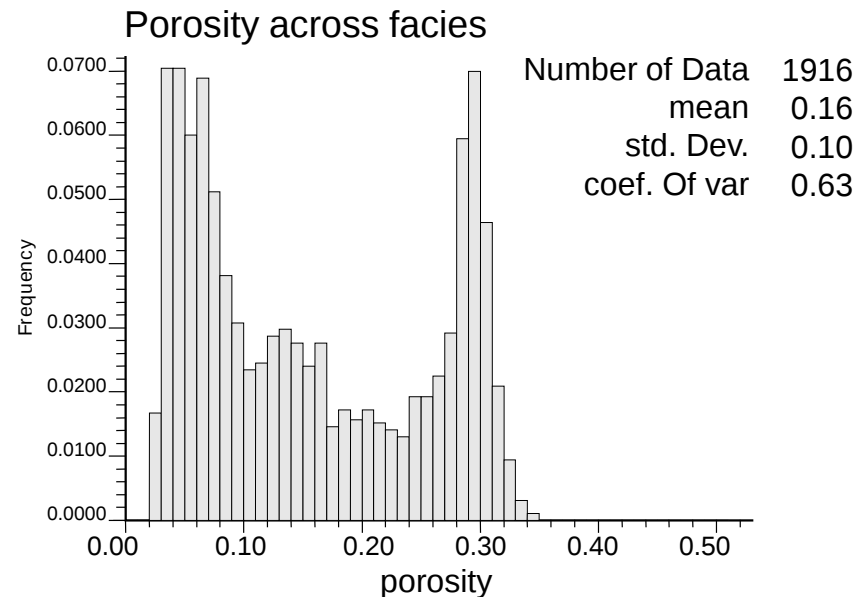
N = 390000
m = 8028.8
s = 1693.7



Impedance

Case Study Summary

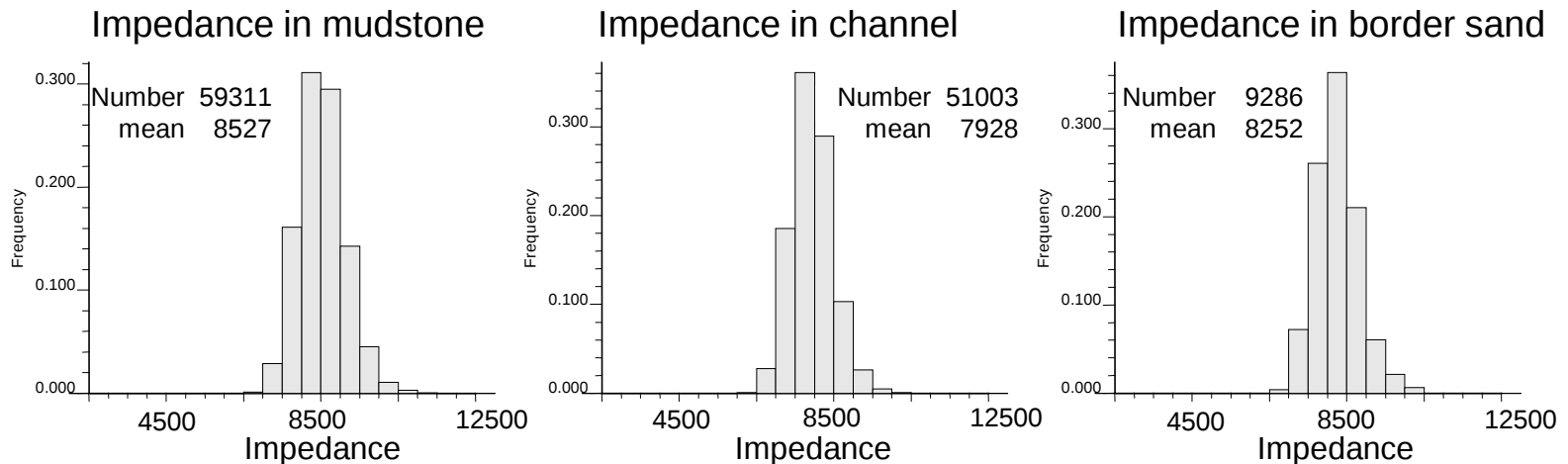
- Reference reservoir is a fluvial reservoir
- Porosity and permeability histogram show distinct bi-modality



Compartmentalization of reservoir into pay / non-pay – decision of stationarity

Facies classification basis

- Seismic impedance in channels, mudstone and border sands :



➡ Pay Facies: channel sands
Non-pay facies: mudstone + splays + levies

Work flow

- simulate **multiple** pay / non-pay facies models
- within non-pay, simulate border facies

➡ **multiple** facies templates

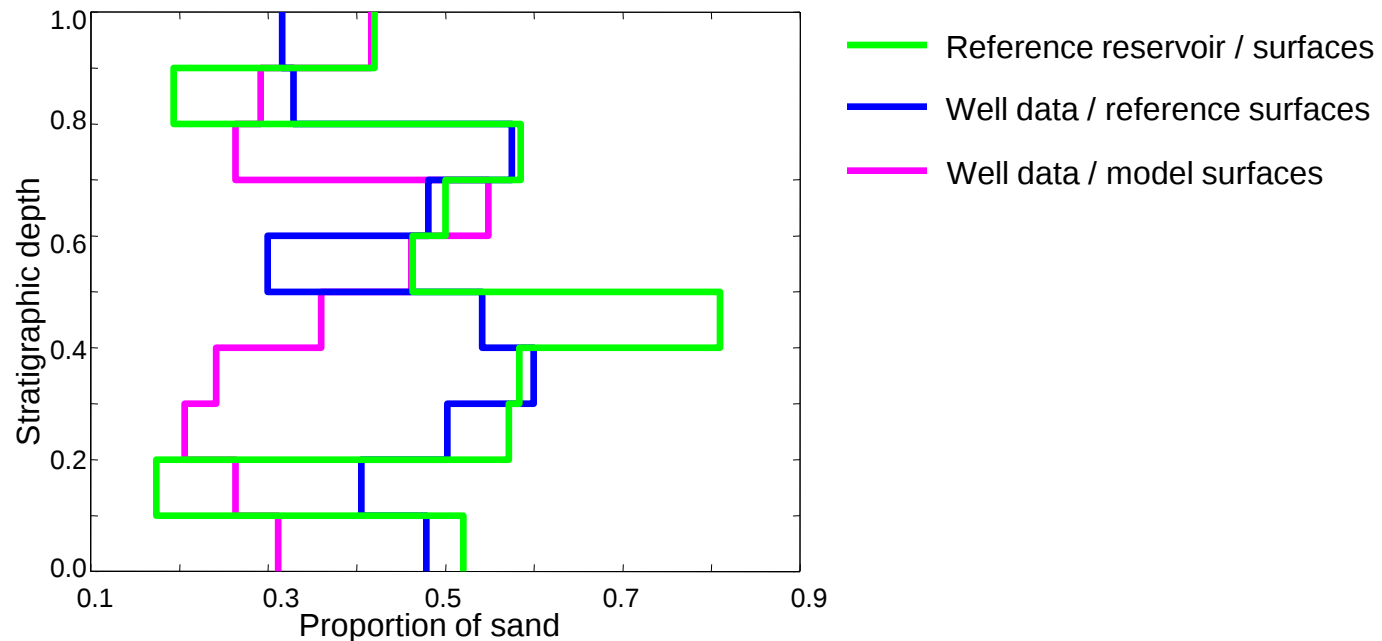
- simulate porosity models separately in pay and non-pay
- same for permeability
- merge pay / nonpay, porosity / permeability models using matching facies template

➡ **multiple** equi-probable reservoir models

Pay / Non-pay facies modeling

Available Data

1) Vertical proportion curve



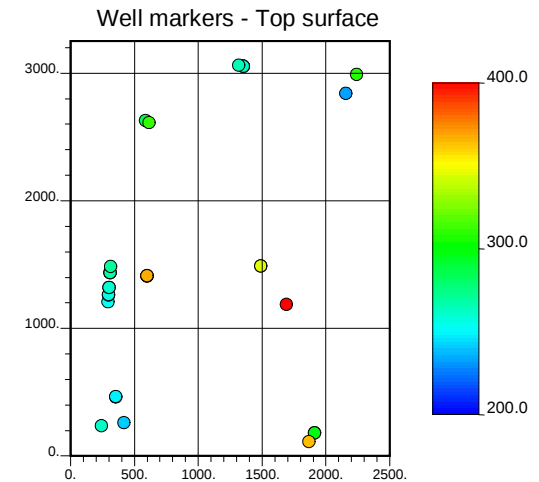
uncertainty in
modeled surfaces



uncertain vertical
proportion curve

Semi-Variogram

Defined as: $\gamma_{XY} = \frac{1}{2} E \left\{ (X - Y)^2 \right\}$



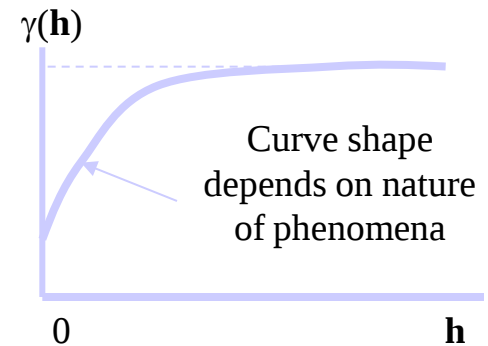
In the spatial context, $2\gamma(\mathbf{u}, \mathbf{u} + \mathbf{h}) = E \left\{ (Z(\mathbf{u}) - Z(\mathbf{u} + \mathbf{h}))^2 \right\}$

Under intrinsic stationarity – the RF $Z(\mathbf{u})$ itself might not be stationary, but increments of the RF $\{Z(\mathbf{u}) - Z(\mathbf{u} + \mathbf{h})\}$ are presumed stationary. Thus:

$$\begin{aligned} \gamma(\mathbf{u}, \mathbf{u} + \mathbf{h}) &= \frac{1}{2} E \left\{ (Z(\mathbf{u}) - Z(\mathbf{u} + \mathbf{h}))^2 \right\} = \frac{1}{2} \text{Var} \{ Z(\mathbf{u}) - Z(\mathbf{u} + \mathbf{h}) \} \\ &= \gamma(\mathbf{h}) \quad \downarrow \text{stationary} \end{aligned}$$

Variogram Inference

- Inference of the experimental variogram requires:
 - Plotting **h**-scatterplot by systematically varying **h** in particular spatial directions.
 - For each scatterplot compute the corresponding moment of inertia of the scatterplot.
 - Plot the moment of inertia versus the lag **h** in specific directions

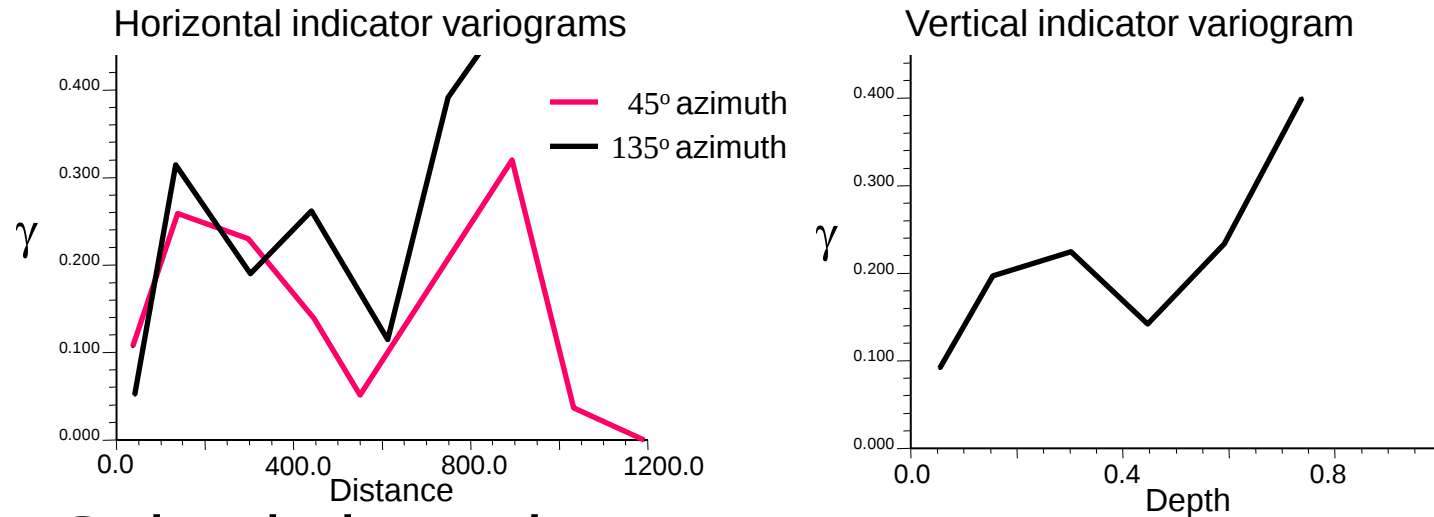


Some remarks

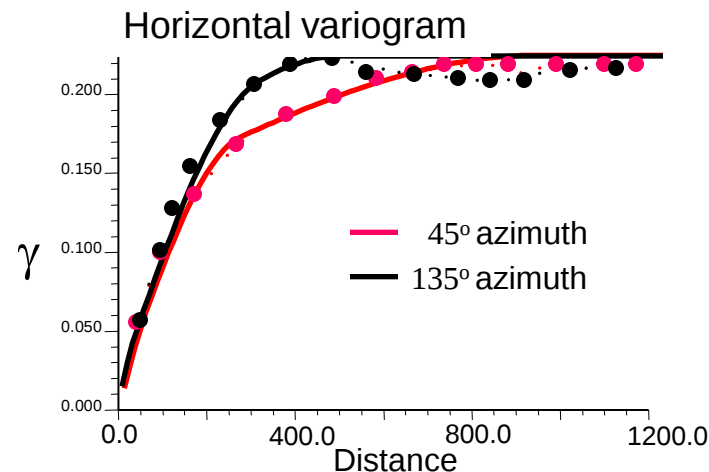
- There must be sufficient number of pairs in each scatterplot (statistical mass)
- Outliers (abnormal high or low values) can cause the moment of inertia to fluctuate wildly

Facies modeling - Data

- Indicator variogram - “hard” data

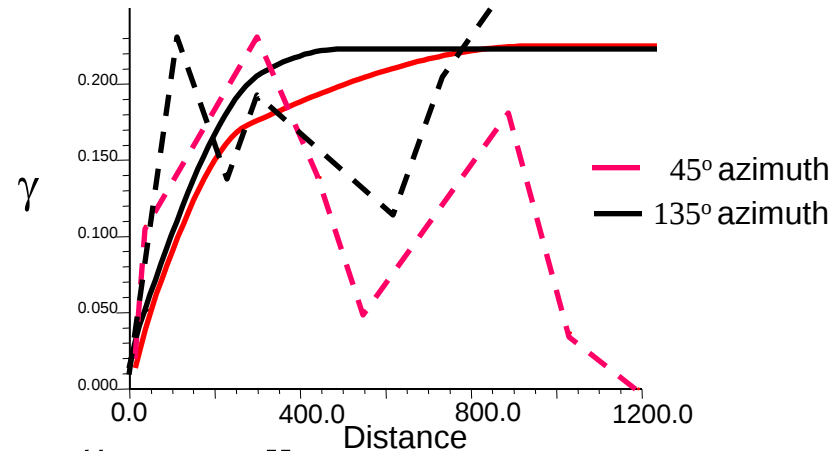


- Seismic impedance

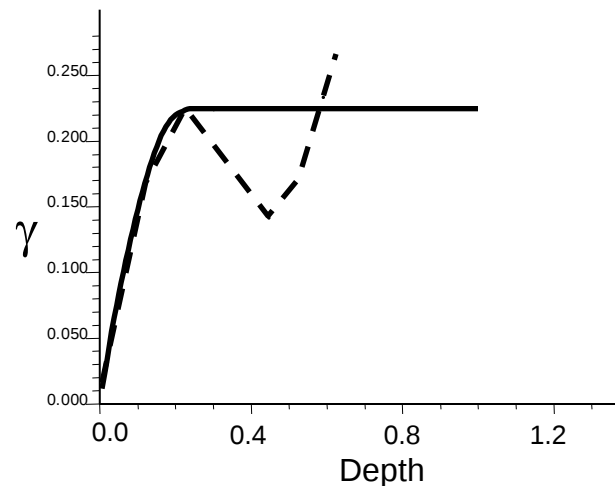


Facies modeling - variograms

Horizontal borrowed from seismic



Vertical from “hard” data



Spatial Estimation

Simple kriging (SK)

of the depth $d(\mathbf{x}_o)$ at unsampled locations $\mathbf{x}_o$
from marker data $d(\mathbf{x}_\alpha), \alpha = 1, \dots, n$

SK estimate:
$$D^*(\mathbf{x}_o) = \left(1 - \sum_{\alpha} \lambda_{\alpha}\right) \cdot m_A + \sum_{\alpha=1}^n \lambda_{\alpha} \cdot D(\mathbf{x}_{\alpha})$$

SK system:
$$\sum_{\beta=1}^n \lambda_{\beta} \cdot C(\mathbf{x}_{\alpha} - \mathbf{x}_{\beta}) = C(\mathbf{x}_{\alpha} - \mathbf{x}_o), \alpha = 1, \dots, n$$

Note: $\mathbf{u} = (\mathbf{x}, \mathbf{d})$

Requisites

$E\{D(\mathbf{x})\} = m_A$ average depth-data over study A

$C(\mathbf{h})$ covariance modeled after a spatial
average $C_A(\mathbf{h})$ over “stationary” area A

Indicator Paradigm

- Consider the indicator RV:

$$I(\mathbf{u}) = \begin{cases} 1, & \text{if } A(\mathbf{u}) = a \\ 0, & \text{otherwise} \end{cases}$$

- Important property:

$$E\{I(\mathbf{u})\} = 1 \times \text{Prob}(A(\mathbf{u}) = a) + 0 \times \text{Prob}(A(\mathbf{u}) \neq a) = \text{Prob}(A(\mathbf{u}) = a)$$

or better still, given n data:

$$E\{I(\mathbf{u}) \mid (n)\} = \text{Prob}(A(\mathbf{u}) = a \mid (n))$$

Indicator basis function

- Consider the expansion $\varphi(n)$ defined on the basis of n indicator random variables

$$: I_{\alpha}, \alpha = 1, \dots, n$$

$$\varphi(n) = a + \sum_{j_1=1}^n b_{j_1}^{(1)} \cdot I(\mathbf{u}_{j_1}) + \sum_{j_1=1}^n \sum_{j_2=1}^n b_{j_1, j_2}^{(2)} \cdot I(\mathbf{u}_{j_1}) I(\mathbf{u}_{j_2}) + \dots + b_{j_1, j_2, \dots, j_n}^{(n)} \prod_{i=1, n} I(\mathbf{u}_i)$$

- Given the n indicator random variables, this expansion is the most complete possible.
- There are a total of $1 + \binom{n}{1} + \binom{n}{2} + \dots + \binom{n}{n} = 2^n$ terms in the above expansion
- There are 2 outcomes for each RV there are
 2^n outcomes possible for the function $\varphi(n)$.
- These outcomes define the space $\varphi(n)$

Indicator Expansion

- The conditional expectation $E\{I(\mathbf{u}) | (n)\}$ is precisely the projection of the unknown indicator event $I(\mathbf{u})$ on to $\varphi(n)$.
- If instead the projection function is defined on a reduced basis L_k :

$$E\{I | (n)\} \cong E_k\{I | (n)\} = I_k^* = a + \sum_{j_1=1}^n b_{j_1}^{(1)} \cdot I(\mathbf{u}_{j_1}) + \sum_{j_1=1}^n \sum_{j_2=1}^n b_{j_1, j_2}^{(2)} \cdot I(\mathbf{u}_{j_1}) I(\mathbf{u}_{j_2}) +$$

$$\square + \sum_{j_1=1}^n \sum_{j_2=1}^n \square \sum_{j_k=1}^n b_{j_1, j_2, \square, j_k}^{(k)} \cdot I(\mathbf{u}_{j_1}) I(\mathbf{u}_{j_2}) \square I(\mathbf{u}_{j_k})$$

Indicator Expansion

- Another way to write the expansion:

$$I_k^* = \sum_{\ell=1}^{n_k} \lambda_{\ell} \cdot V_{\ell} \quad n_k = 1 + \binom{n}{1} + \binom{n}{2} + \cdots + \binom{n}{k}$$

basis function $V_1 = \{\mathbf{1}\}$; $\{I_i\}$, $\{I_i \cdot I_j\}$ etc.

Projection Theorem

Let H be a Hilbert space and M a closed subspace of H , then corresponding to any $x \in H$ there exists a unique vector $m_o \in M$ such that:

$$\|x - m_o\| \leq \|x - m\| \quad \forall m \in M$$

Furthermore a necessary and sufficient condition that $m_o \in M$ be a unique minimizing vector is that $x - m_o$ be orthogonal to M (Luenberger 1997, p. 51).

Normal Equations

- The implication of the projection theorem is that:

$$\left\langle (I - I_k^*), V_\ell \right\rangle = 0 \quad \Leftrightarrow \quad \langle I \cdot V_\ell \rangle = \langle I_k^* \cdot V_\ell \rangle$$

or in terms of projections:

$$E\{I_k^* V_\ell\} = E\{I V_\ell\} , \quad \forall \ell = 1, \dots, n_k$$

which is a system of n_k normal equations

Examples

$$V_1 = \mathbf{1} \Rightarrow E\{I_k^* \mathbf{1}\} = \sum_{\ell=1}^{n_k} \lambda_{\ell} \cdot E\{V_{\ell}\} = E\{I\} \quad \text{- unbiasedness}$$

$$V_{\ell} = I_j, \quad j = 1, \dots, n \Rightarrow E\{I_k^* I_j\} = \sum_{\ell=1}^{n_k} \lambda_{\ell} \cdot E\{V_{\ell} \cdot I_j\} = E\{I \cdot I_j\}$$

- reproduction of indicator covariances

$$V_{\ell} = I_{j_1} \cdot I_{j_2}, \quad j_1, j_2 = 1, \dots, n \Rightarrow E\{I_k^* I_{j_1} I_{j_2}\} = \sum_{\ell=1}^{n_k} \lambda_{\ell} \cdot E\{V_{\ell} \cdot I_{j_1} I_{j_2}\} = E\{I \cdot I_{j_1} I_{j_2}\}$$

- reproduction of the 3rd order covariances

In general:

$$\sum_{\ell=1}^{n_k} \lambda_{\ell} \cdot E\{V_{\ell} \cdot V'_{\ell}\} = E\{I \cdot V'_{\ell}\}, \quad \ell = 1, \dots, n_k$$

requires knowledge of up to order $k + 1$ indicator covariances

Facies modeling - indicator approach

Indicator kriging (Simple IK)

$$I^*(\mathbf{u}) = p + \sum_{\alpha=1}^{n(\mathbf{u})} \lambda(\mathbf{u}_{\alpha}) \cdot [I(\mathbf{u}_{\alpha}) - p]$$

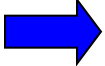
Interpretation

$$I^*(\mathbf{u}) = Prob\{u \in \text{pay} | n \text{ data}\}$$

IK System

$$\sum_{\beta=1}^{n(\mathbf{u})} \lambda(\mathbf{u}_{\beta}) \cdot C_I(\mathbf{h}_{\alpha\beta}) = C_I(\mathbf{h}_{\alpha 0})$$

Facies modeling - indicator approach

Indicator kriging  local conditional probability
 $Prob\{u \in \text{pay} | n \text{ data}\}$

Simulation

- Draw $r \in \text{Uniform}[0,1]$
If $r \leq Prob\{u \in \text{pay} | n \text{ data}\}$
then $I^\ell(\mathbf{u}) = 1: \mathbf{u} \in \text{pay}$ else $I^\ell(\mathbf{u}) = 0: \mathbf{u} \in \text{nonpay}$
- Add simulated $I^\ell(\mathbf{u})$ into data set : $(n) \rightarrow (n+1)$
- Visit next node $\mathbf{u}'$

Facies modeling - indicator approach

Locally varying mean

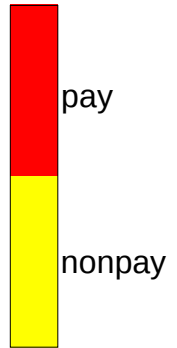
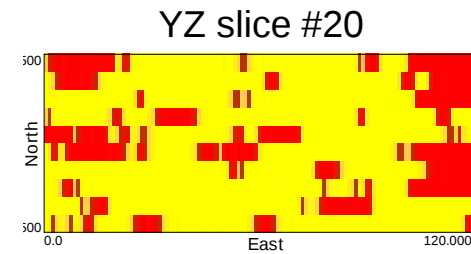
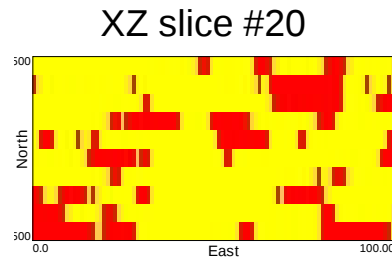
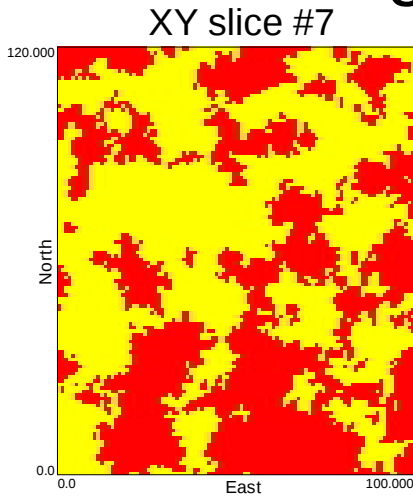
$$I^*(\mathbf{u}) = p_V(d) + \sum_{\alpha=1}^{n(\mathbf{u})} \lambda(\mathbf{u}_\alpha) \cdot [I(\mathbf{u}_\alpha) - p_V(d_\alpha)]$$

Requisites

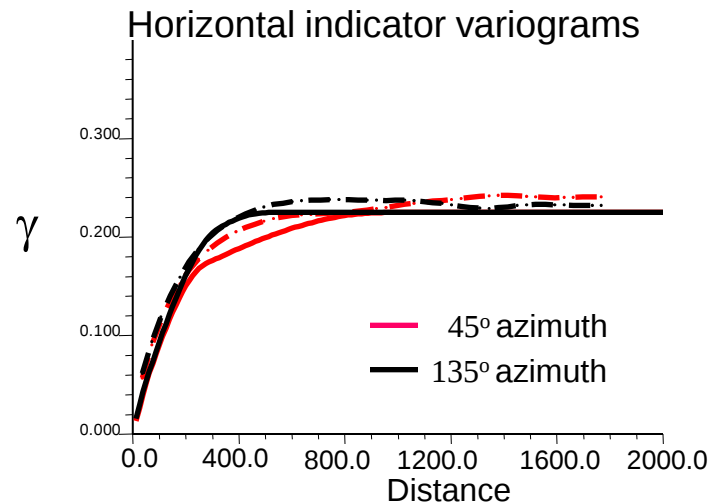
- vertical proportion of pay facies in stratigraphic layer d : $p_V(d)$
- Indicator variogram model, $\gamma_I(\mathbf{h})$

Indicator Simulation - Results

- Slices through the facies model

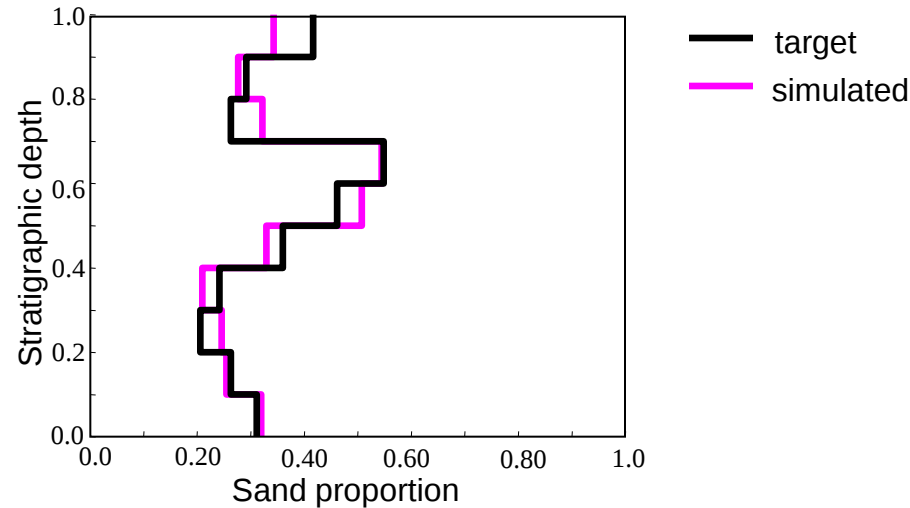


- Variogram reproduction

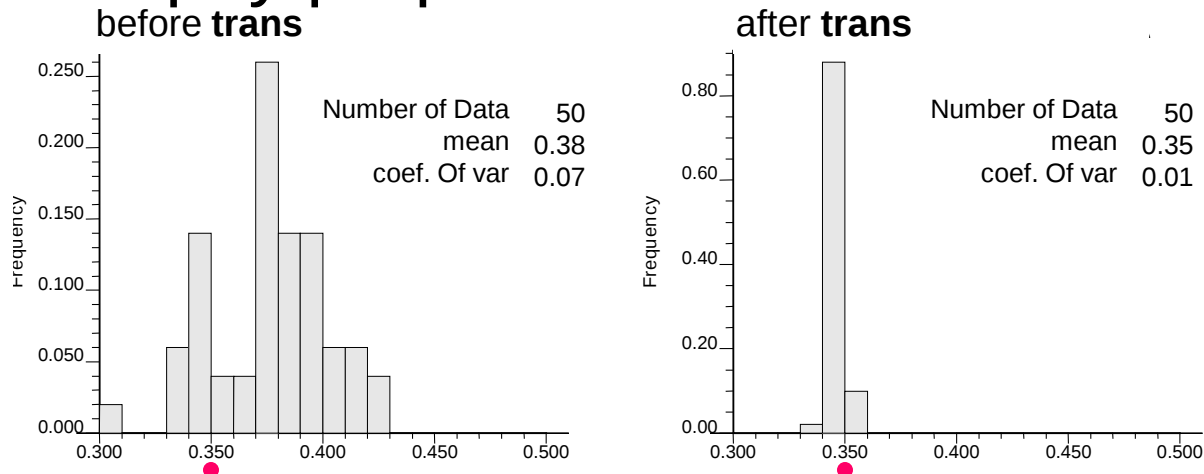


Indicator Simulation - Results

- Vertical proportion

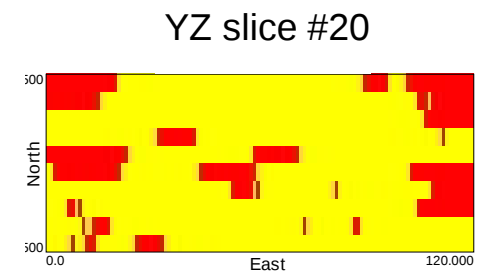
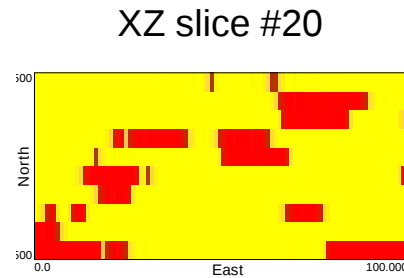
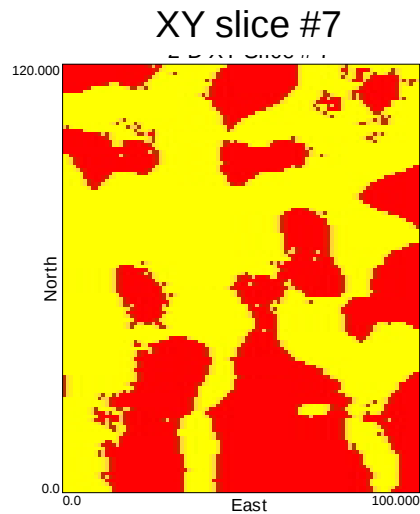


- Global pay proportion



Indicator Simulation - Results

- Perform a histogram transformation
- Simulation after **trans**



Facies modeling - using seismic info.

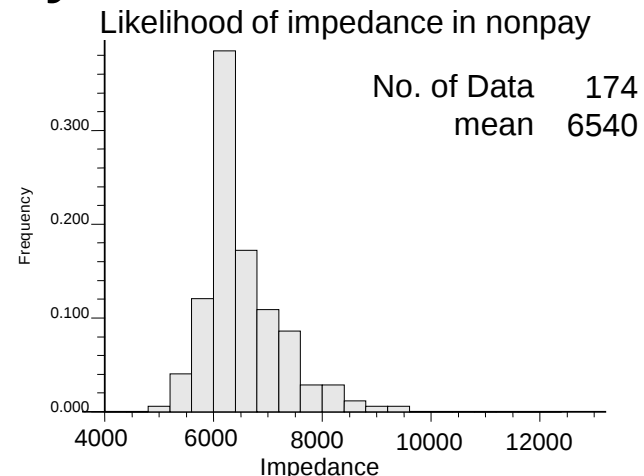
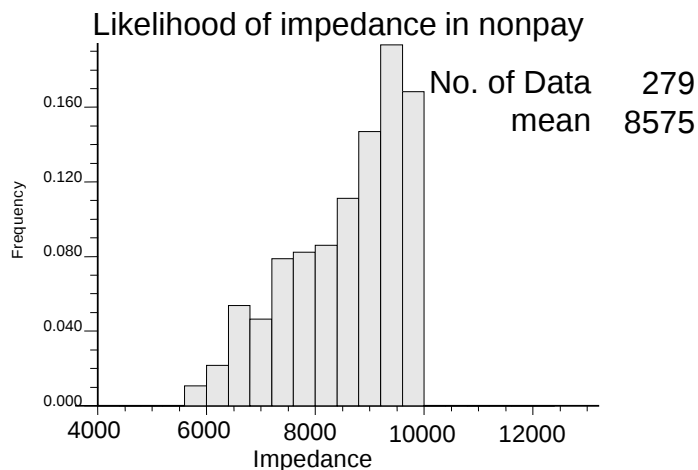
- Likelihood of seismic impedance | $u \in \text{pay}$

$$F(s | 1) = \text{Prob}\{S(\mathbf{u}) \leq s \mid \mathbf{u} \in \text{pay}\}$$

Similarly,

$$F(s | 0) = \text{Prob}\{S(\mathbf{u}) \leq s \mid \mathbf{u} \in \text{nonpay}\}$$

- If: $F(s | 1) \neq F(s | 0)$ then impedance discriminates pay / nonpay



Facies modeling - using seismic info.

- Want $Prob\{I(\mathbf{u})=1 \mid s(\mathbf{u}) \in [s_l, s_{l+1}]\}$

- Apply Bayes' inversion:

$$Prob\{A \mid B\} = \frac{Prob\{B \mid A\}}{Prob\{B\}} \cdot Prob\{A\}$$

$$\begin{aligned} & Prob\{I(\mathbf{u})=1 \mid s(\mathbf{u}) \in [s_l, s_{l+1}]\} \\ &= \frac{Prob\{s(\mathbf{u}) \in [s_l, s_{l+1}] \mid I(\mathbf{u})=1\}}{Prob\{s(\mathbf{u}) \in [s_l, s_{l+1}]\}} \cdot Prob\{I(\mathbf{u})=1\} \\ &= \frac{F(s_{l+1} \mid 1) - F(s_l \mid 1)}{Prob\{s(\mathbf{u}) \in [s_l, s_{l+1}]\}} \cdot p(\mathbf{u}) \end{aligned}$$

Facies modeling - Bayes' inversion

Requisites

- Prior vertical proportion, $p(d)$
- Likelihood, $F(s | 1)$
- Seismic impedance distribution, $F(s)$

Implementation

$$p'(\mathbf{u}) = \text{Prob}\{I(\mathbf{u})=1 \mid s(\mathbf{u}) \in [s_l, s_{l+1}]\}$$

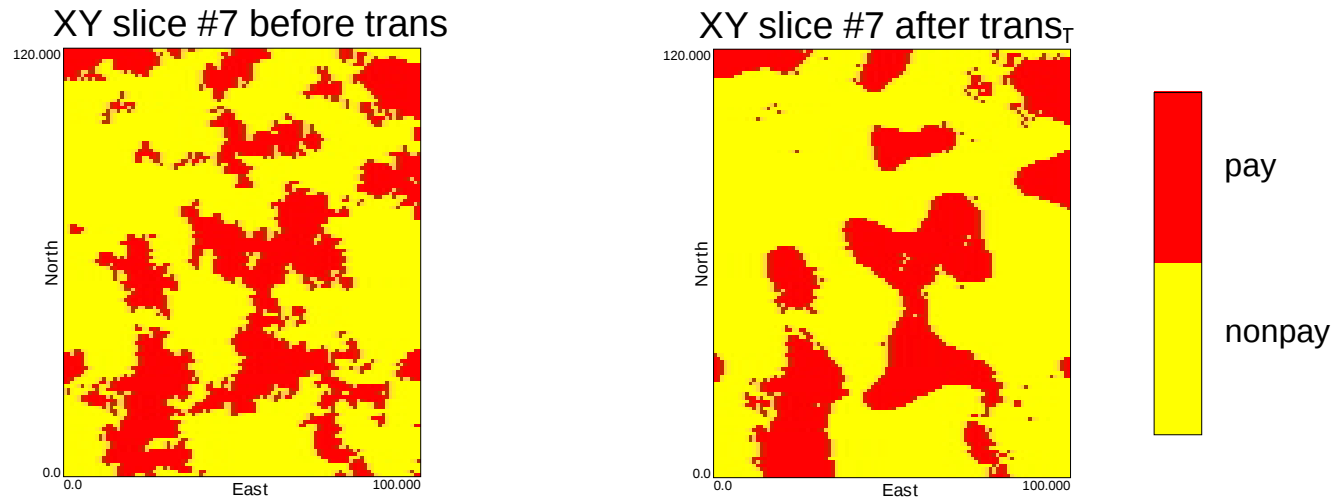
used as locally varying mean after standardization

$$\frac{1}{N_d} \sum_x p'(x, d) = \underbrace{p_V(d)}_{\text{layer d pay proportion}} \text{ where } \mathbf{u}=(\mathbf{x}, \mathbf{d})$$

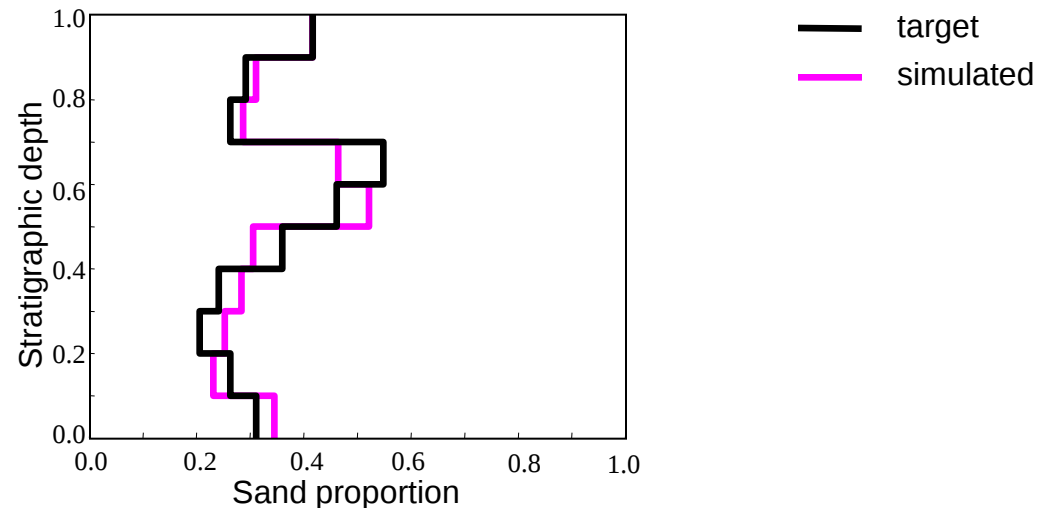
layer d pay proportion

Bayes' inversion - Results

- Slices through facies model



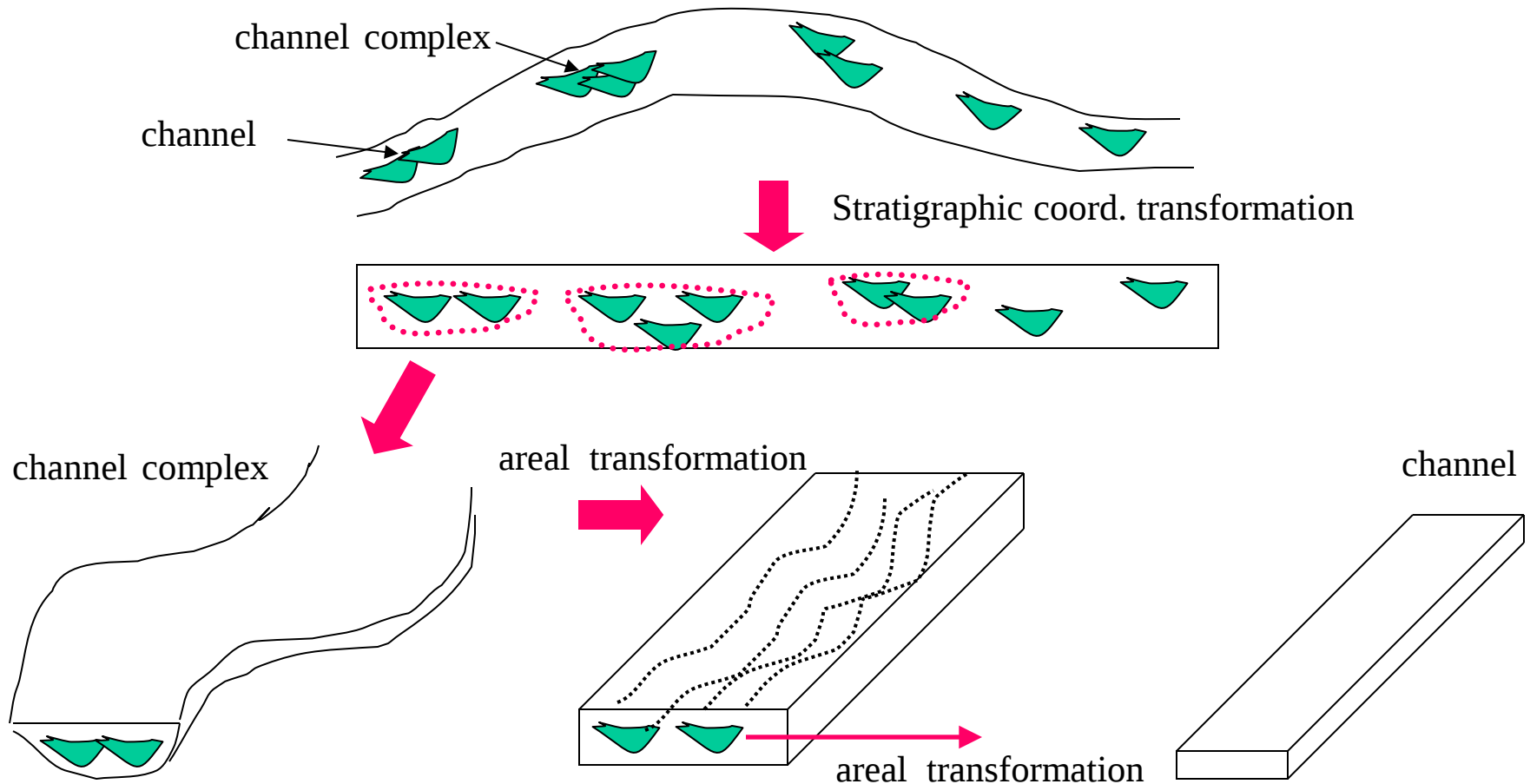
- Vertical proportion



Facies modeling- object based

fluvsim (Deutsch and Wang (1996))

Modeling using reversible coord. transforms



Facies modeling - object based

Simulation procedure in transformed space

- annealing simulation of channel complex

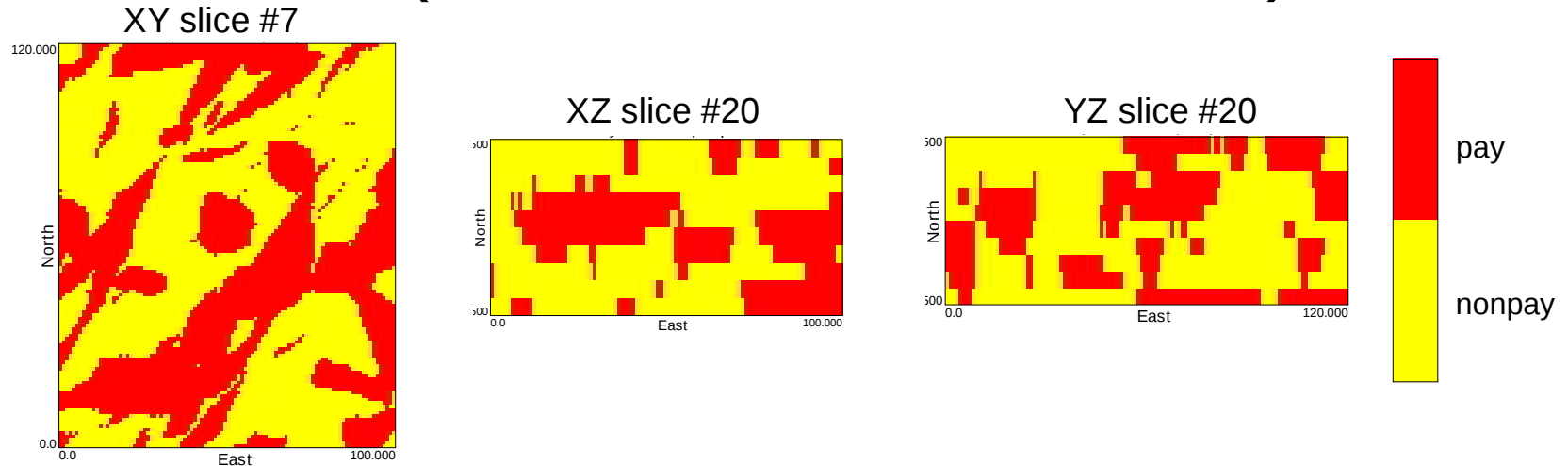
$$O_{cc} = \underbrace{|p_G - p_G^*|}_{\text{global proportion}} + \sum_{n_z} \underbrace{|p_V(d) - p_V(d)^*|}_{\text{vertical proportion}} + \sum_n \underbrace{|I(\mathbf{u}_\alpha) - I_{CC}(\mathbf{u}_\alpha)|}_{\text{well data}}$$

where $I(\mathbf{u})=1$ if simulated $\mathbf{u} \in$ channel

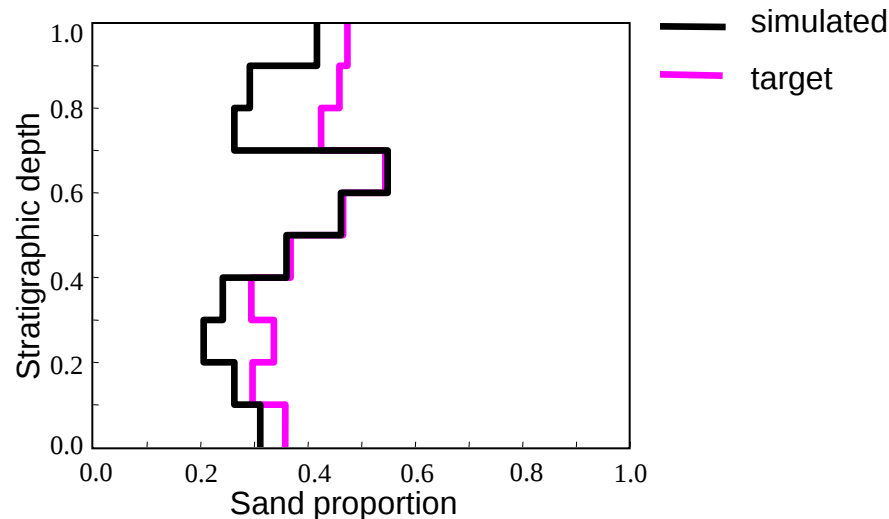
- perturb channel complex until convergence
- annealing simulation of channel
- perturb channel until convergence
- back coord. transform to obtain reservoir model

Fluvsim - results

Facies model (well mismatch: 11.2 %)

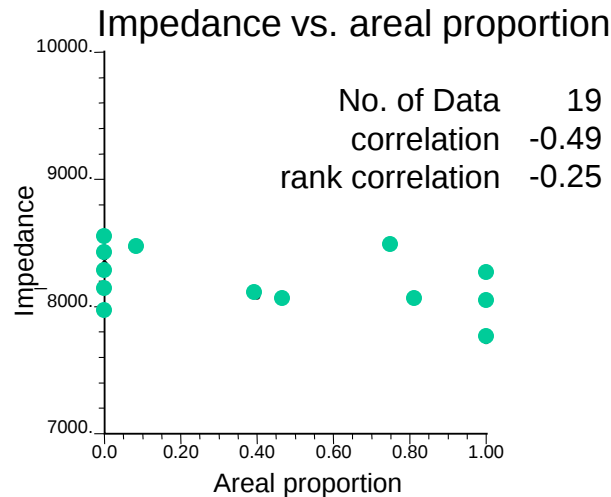


Vertical proportion curve



Fluvsim model - using seismic info.

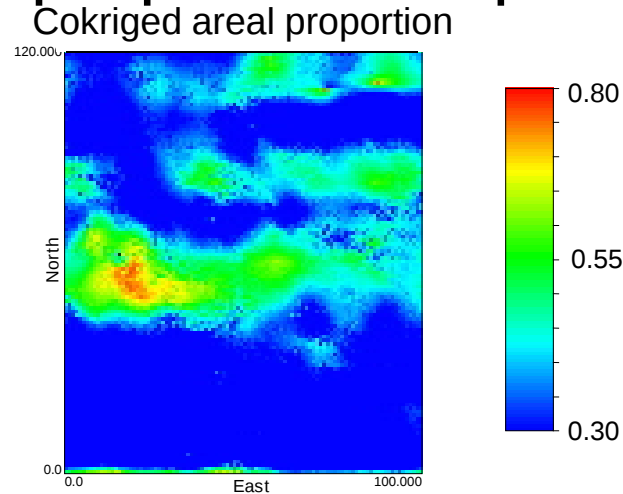
- Average impedance vs. areal pay proportion



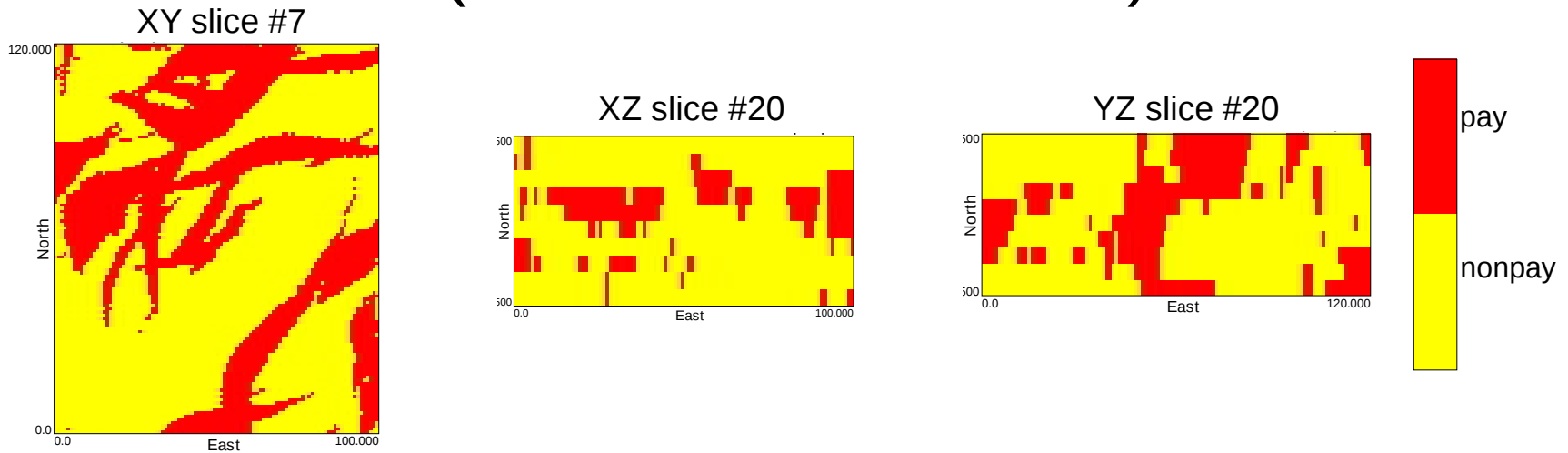
- ➡ Cokrige areal proportion map using:
- vertically averaged proportion data
 - exhaustive 2-D seismic impedance
 - variogram of facies proportion
 - co-located cokriging under MMI

Fluvsim - using seismic info.

- Cokriged areal proportion map

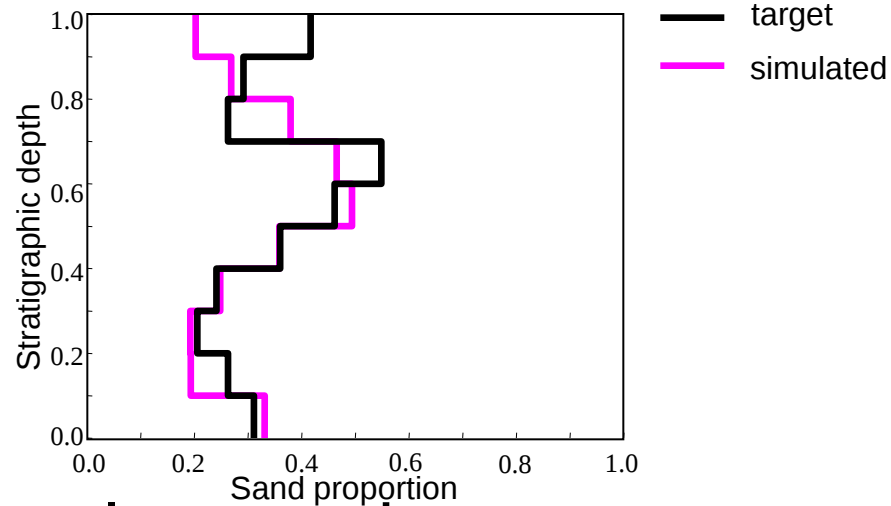


- **fluvsim** slices (well mismatch - 14 %)



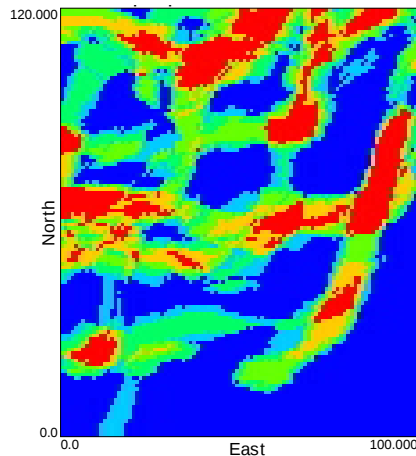
Fluvsim - using seismic info.

- fluvsim vertical proportion

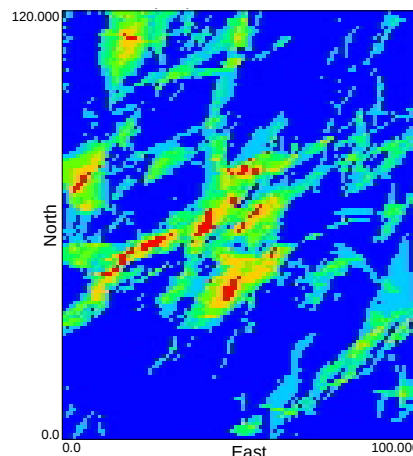


- impact of areal proportion

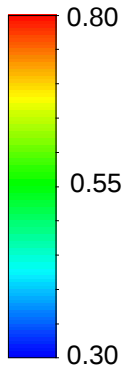
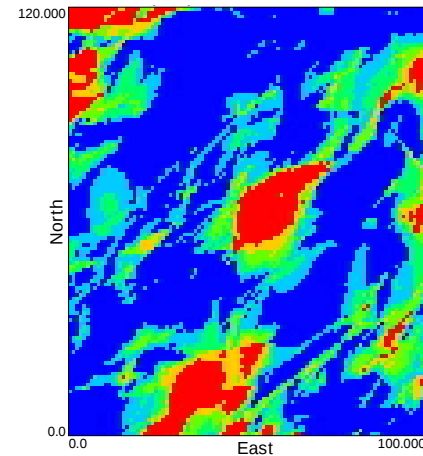
Areal proportion - reference



Areal proportion - reproduction



Areal prop. - unconditional



Facies modeling - Some remarks

- Object based models geologically realistic **but** require numerous shape parameters
- **trans** post-processing of indicator simulations essential
- Integration of seismic has significant impact

Indicator basis function

- Consider the expansion $\varphi(n)$ defined on the basis of n indicator random variables

$$: I_{\alpha}, \alpha = 1, \dots, n$$

$$\varphi(n) = a + \sum_{j_1=1}^n b_{j_1}^{(1)} \cdot I(\mathbf{u}_{j_1}) + \sum_{j_1=1}^n \sum_{j_2=1}^n b_{j_1, j_2}^{(2)} \cdot I(\mathbf{u}_{j_1}) I(\mathbf{u}_{j_2}) + \dots + b_{j_1, j_2, \dots, j_n}^{(n)} \prod_{i=1, n} I(\mathbf{u}_i)$$

- Given the n indicator random variables, this expansion is the most complete possible.
- There are a total of $1 + \binom{n}{1} + \binom{n}{2} + \dots + \binom{n}{n} = 2^n$ terms in the above expansion
- There are 2 outcomes for each RV there are
 2^n outcomes possible for the function $\varphi(n)$.
- These outcomes define the space $\varphi(n)$

Indicator Expansion – Another Interpretation

In general the extended equations are for all 2^n realizations of the n data.

If instead, the estimate I_o^* corresponding to a specific realization of the indicator RVs, say

$$D = \{i_1, i_2, \dots, i_n\} = 1, \text{ then: } I_o^* = \varphi(D) = \lambda_o + \lambda_1 \cdot D$$

Two bases $1, D$ – two equations to obtain two unknowns

Unbiasedness: $I_o^* - p_o = \lambda_1 \cdot (D - E\{D\})$

Orthogonality:

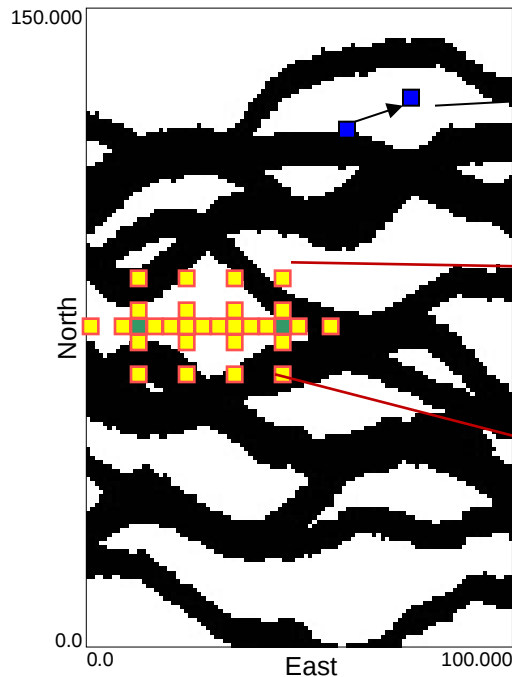
$$\langle I_o - I_o^*, D \rangle = \langle I_o - p_o - \lambda_1 \cdot (D - E\{D\}), D \rangle = 0$$

$$\Rightarrow \lambda_1 = \frac{E(I_o D) - p_o E(D)}{\text{Var}\{D\}} = \frac{E(I_o D) - E(I_o)E(D)}{\text{Var}\{D\}}$$

Single Normal Equation

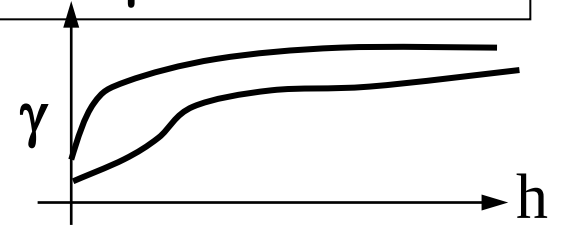
Calibration of geologic information

Geologic model



traditional

Variogram – two-point statistics



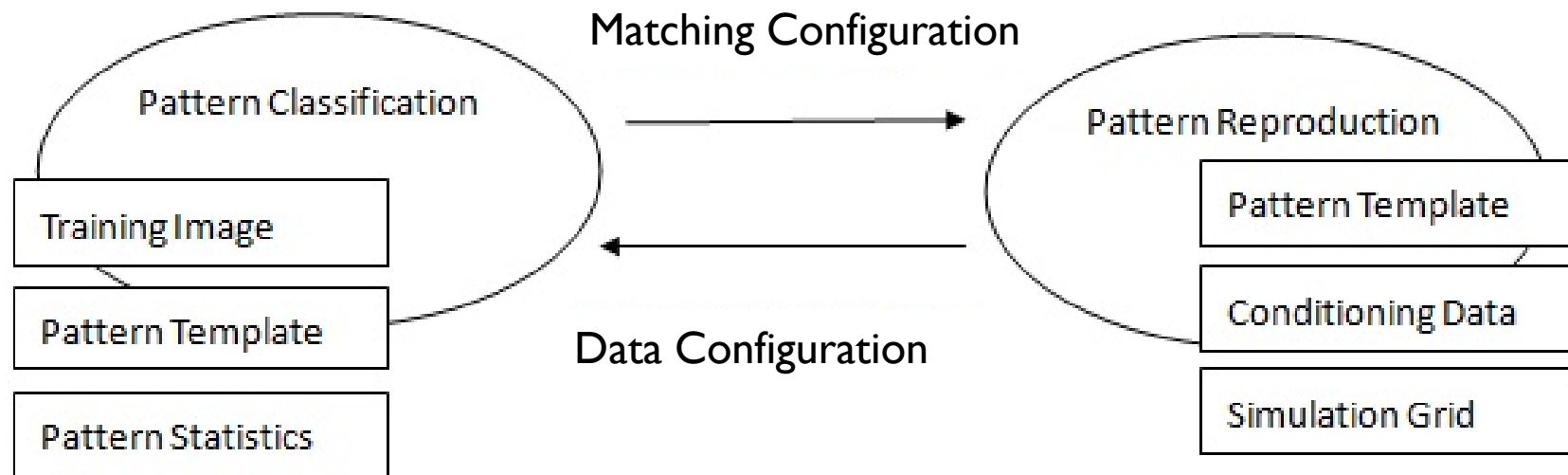
$i(u; s_k)$

mp statistics $s_t(u) = \{ s(u+h_\alpha), \alpha = 1, \dots, n_t \}$

Set of data events = training data set

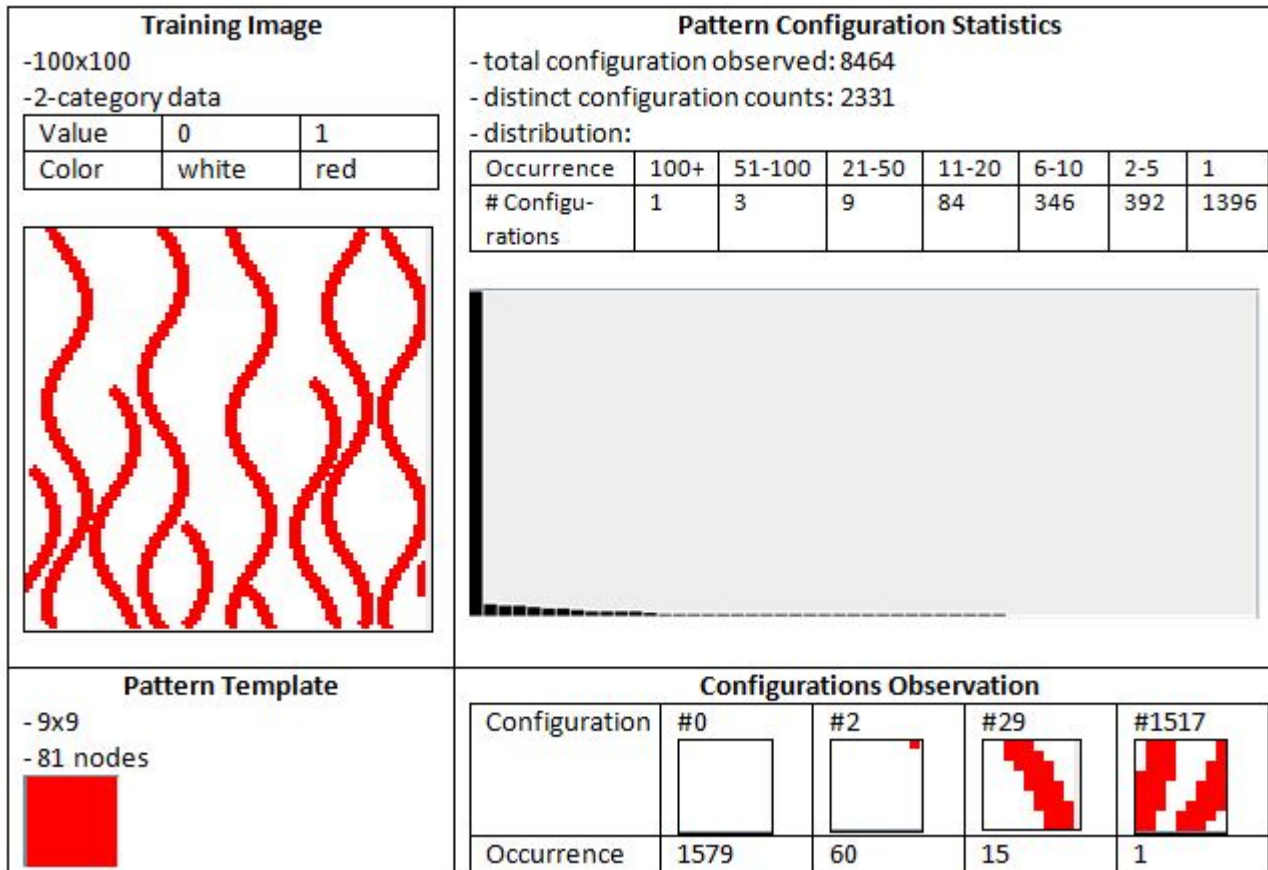
GrowthSim

- components



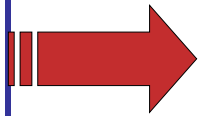
Pattern Statistics

- pattern observation



Multipoint Statistics

To condition realization to hard (static) data



Sample from distribution $P(A | B)$

Geological / Hard information
Well data

Use Bayes' Rule

$$P(A | B) = \frac{P(A, B)}{P(B)}$$

MP Histogram

where A = Simulation pattern

B = Information from conditioning data & geology

A – single point simulation event

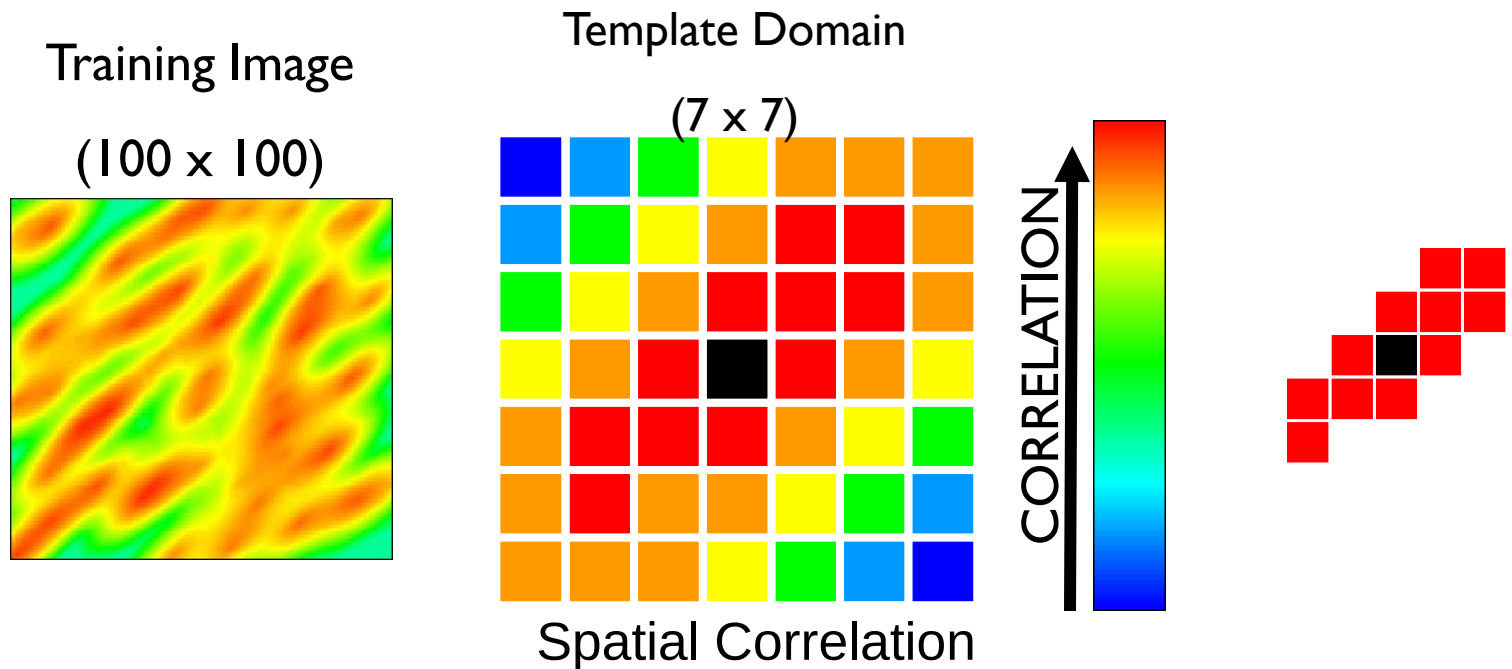
Strebelle (2000), Caers (1999)

A – multiple point simulation event

GROWTHSIM, Arpat (2005)

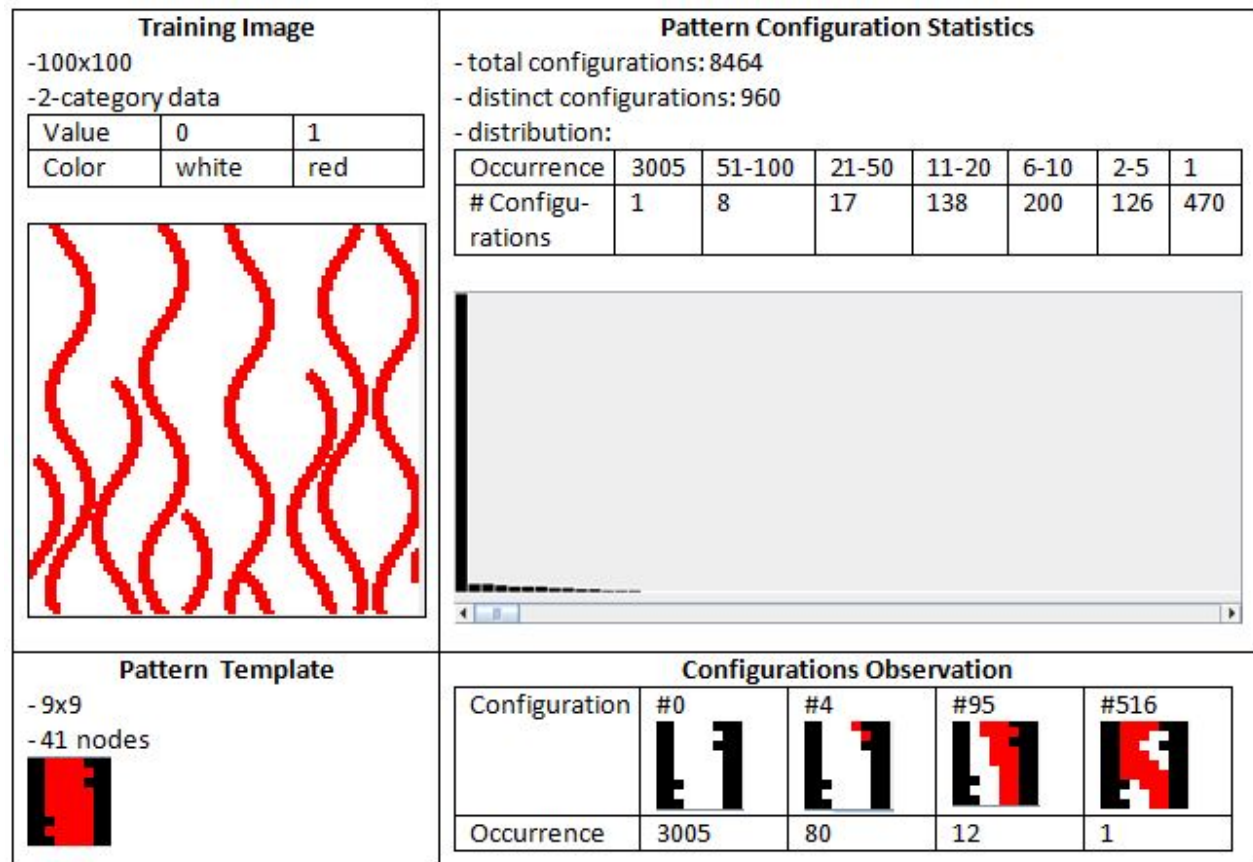
Pattern Statistics

- optimized template
 - goal
 - reduce noises
 - reduce spurious pattern configurations



Pattern Statistics

- effects of optimized template

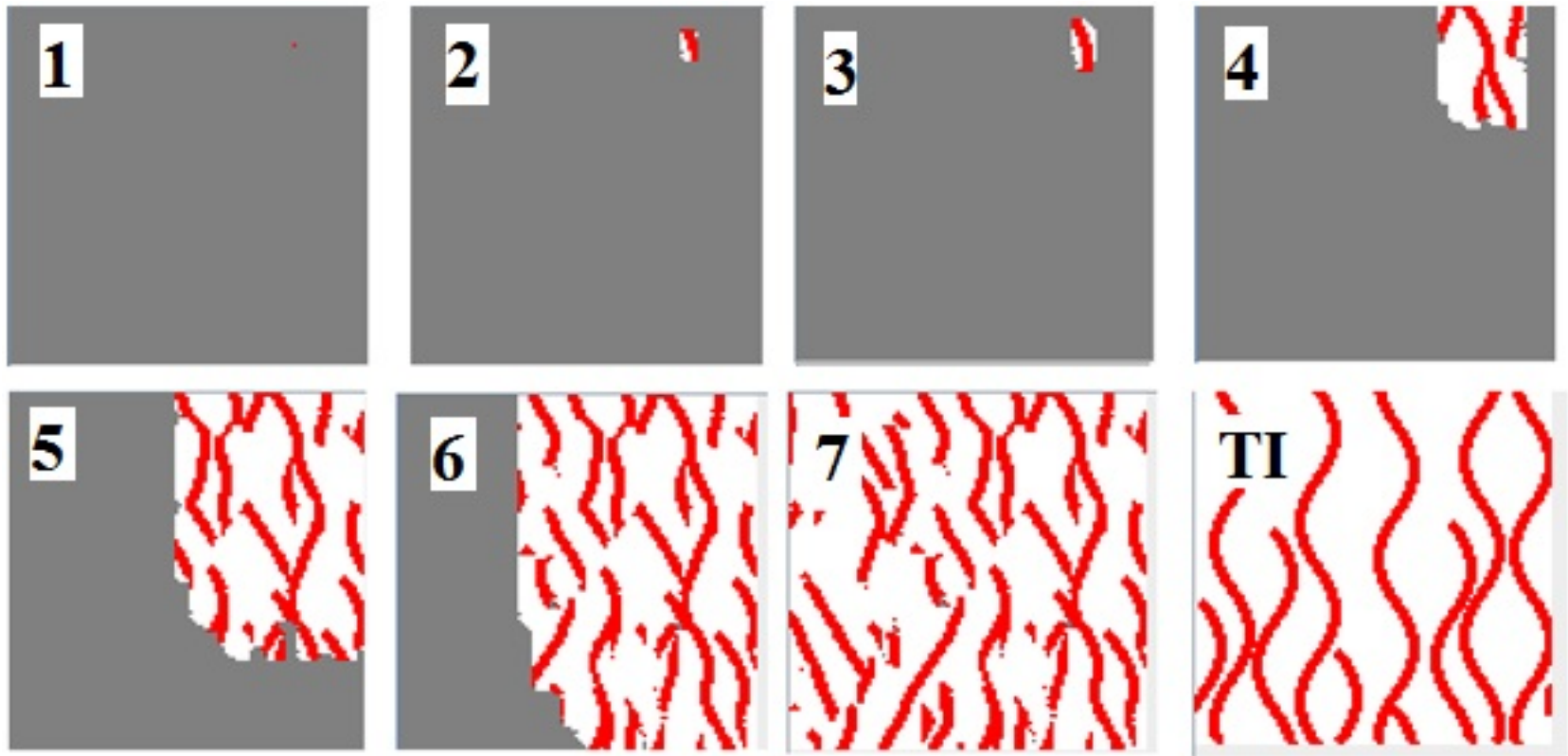


GrowthSim

- algorithm
 1. gather conditioning data around some node in the simulation grid
 2. look for matching pattern configurations to the conditioning data in the statistics
 3. apply one of the matching pattern configuration to the simulation grid
 4. repeat 1-3 around the newly simulated node, until the simulation grid is filled

GrowthSim

- 100x100, 2-category data TI

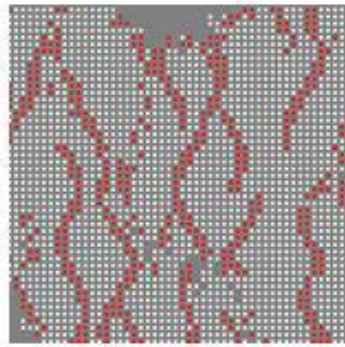


Multiple Grid Simulation

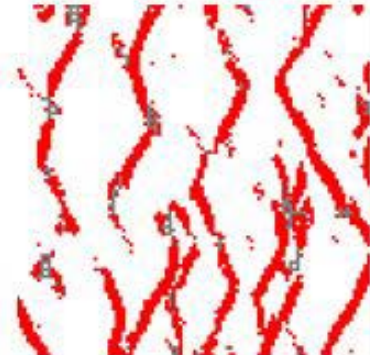
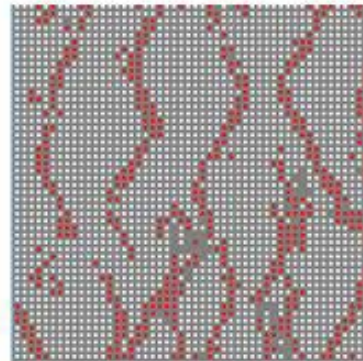
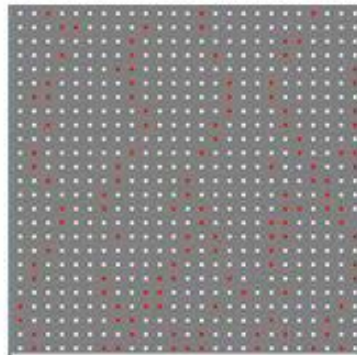
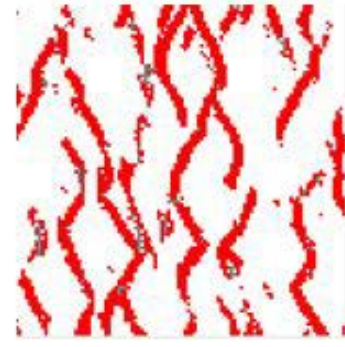
- effects of multiple-level simulation



(a) Using level 1 only



(b) Using level 2 and 1



(c) Using level 3, 2, and 1

Calibration of geologic information

An exhaustive training image required for inference of statistics

Noisy spatial covariance (variogram) inferred using sparse data,
modeled using smooth functions → restrictive, unrealistic

- Develop a geostatistical simulation method that does not rely on an exhaustive training model
- Perform inference and modeling of spatial connectivity functions in the spectral domain using sparse data.

Moments and Cumulants

Moment Generating Function

$$M[\omega] = E[e^{\omega z}] = \int_{-\infty}^{\infty} e^{\omega z} f_Z(z) dz$$

Expanding the function $e^{\omega z}$ as a Taylor series about the origin:

$$\begin{aligned} M[\omega] &= E[e^{\omega z}] = E \left[\sum_{r=0}^{\infty} \frac{\omega^r z^r}{r!} \right] \\ &= \sum_{r=0}^{\infty} \frac{\omega^r E[z^r]}{r!} = \sum_{r=0}^{\infty} \frac{\omega^r \text{Mom}[z, \dots, z]}{r!} \end{aligned}$$

Cumulant generating function of Z is the Neperian logarithm of the moments generating function M :

$$\begin{aligned} K[\omega] &= \ln(E[e^{\omega z}]) = \ln \left(E \left[\sum_{r=0}^{\infty} \frac{\omega^r z^r}{r!} \right] \right) \\ &= \sum_{r=0}^{\infty} \frac{\omega^r \text{Cum}[z, \dots, z]}{r!} \end{aligned}$$

Moments and Cumulants

Relation between the first few moments and cumulants

$$\text{Cum}[z] = \text{Mom}[z]$$

$$\text{Cum}[z, z] = \text{Mom}[z, z] - \text{Mom}[z]^2$$

$$\text{Cum}[z, z, z] = \text{Mom}[z, z, z] - 3\text{Mom}[z, z] \cdot \text{Mom}[z] + 2\text{Mom}[z]^3$$

Moments can be calculated from the cumulants by

$$\text{Mom}[z, z] = \text{Cum}[z, z] + \text{Cum}[z]^2$$

$$\text{Mom}[z, z, z] = \text{Cum}[z, z, z] + 3\text{Cum}[z, z]\text{Mom}[z] + \text{Cum}[z]^3$$

Three-point moment is a measure of similarity between three spatial locations. High value implies three locations jointly have high values of the spatial attribute Z **Spatial cumulant is also a measure of spatial connectivity.**

Polyspectra

Assuming $C_{k,z}(\mathbf{h}_1, \mathbf{h}_2, \dots, \mathbf{h}_{k-1})$ is absolutely summable (i.e. $\sum_{\mathbf{h}_1=-\infty}^{\infty} \dots \sum_{\mathbf{h}_{k-1}=-\infty}^{\infty} |C_{k,z}(\mathbf{h}_1, \mathbf{h}_2, \dots, \mathbf{h}_{k-1})|$ exists and is finite), the k -th order polyspectrum is defined as the $(k-1)$ -dimensional discrete Fourier transform of the k -th order cumulant, i.e.,

$$S_{k,z}(\omega_1, \omega_2, \dots, \omega_{k-1}) = \sum_{\mathbf{h}_1=-\infty}^{\infty} \dots \sum_{\mathbf{h}_{k-1}=-\infty}^{\infty} C_{k,z}(\mathbf{h}_1, \mathbf{h}_2, \dots, \mathbf{h}_{k-1}) \times \exp\left[-j \sum_{i=1}^{k-1} \omega_i \mathbf{h}_i\right]$$

$S_{2,z}(\omega)$ is the Fourier transform of covariance function  power spectrum

$S_{3,z}(\omega_1, \omega_2)$: bispectrum (also noted as $B(\omega_1, \omega_2)$)

$S_{4,z}(\omega_1, \omega_2, \omega_3)$: trispectrum (also noted as $T(\omega_1, \omega_2, \omega_3)$)

Power Spectrum

The power spectrum is the Fourier transform of the covariance function - a measure of the energy of a signal. For a deterministic signal $z(t)$, energy:

$$E = \int_{-\infty}^{\infty} |z(t)|^2 dt = \int_{-\infty}^{\infty} P(\omega) d\omega \quad P(\omega) \text{ is the power spectrum of the signal}$$

$$Z(\omega) = \int_{-\infty}^{\infty} z(t) e^{-i\omega t} dt \text{ and the inverse } z(t) = \int_{-\infty}^{\infty} Z(\omega) e^{i\omega t} d\omega$$

$$P(\omega) = Z(\omega) Z^*(\omega) = |Z(\omega)|^2 \quad Z^*(\omega) \text{ -conjugate of Fourier transform } Z(\omega)$$

For a stochastic process $Z(t)$, energy is defined as

$$E = E\{|Z(t)|^2\} = \int_{-\infty}^{\infty} |Z(t)|^2 dz = \int_{-\infty}^{\infty} P(\omega) d\omega$$

$$P(\omega) = \lim_{T \rightarrow \infty} E\{Z_T(\omega) Z_T^*(\omega)\} = \lim_{T \rightarrow \infty} E\{|Z_T(\omega)|^2\} \quad Z_T(\omega) \text{ - Fourier transform of } z_T(t)$$

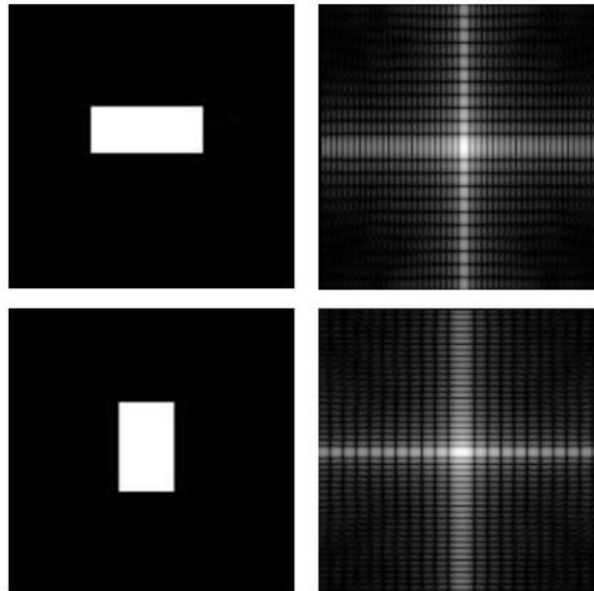
Power Spectrum – Detecting orientations

Property of Fourier transform - if an image is rotated, its Fourier transform also rotates

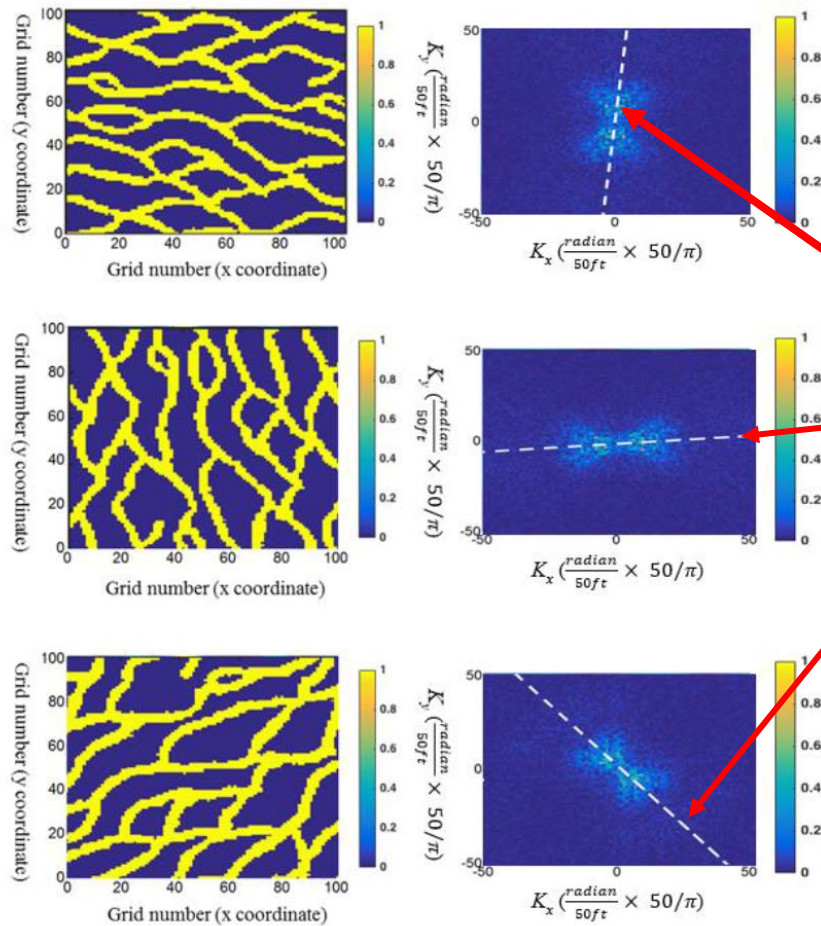
changes in direction of spatial continuity  rotation in Fourier transform space



orientation of objects can be detected from power spectrum



Power Spectrum – Detecting orientations



Orientation calculated using eigenvectors of the inertia matrix computed by thresholding the Fourier transform

Bispectrum

For a deterministic signal $z(t)$:
$$\int_{-\infty}^{\infty} |z(t)|^3 dt = \int_{-\infty}^{\infty} \int_{-\infty}^{\infty} B(\omega_1, \omega_2) d\omega_1 d\omega_2$$

$$B(\omega_1, \omega_2) = Z(\omega_1)Z(\omega_2)Z^*(\omega_1 + \omega_2) \quad \text{Bispectrum}$$

For a stationary random process :
$$E\{|z(t)|^3\} = \int_{-\infty}^{\infty} \int_{-\infty}^{\infty} B(\omega_1, \omega_2) d\omega_1 d\omega_2$$

$$B(\omega_1, \omega_2) = \lim_{T \rightarrow \infty} E\{Z_T(\omega_1)Z_T(\omega_2)Z_T^*(\omega_1 + \omega_2)\} \quad \text{Bispectrum}$$

Power spectrum contribution of each individual frequency component independently,

Bispectrum includes frequency interactions between components ω_1 , ω_2 , and $\omega_1 + \omega_2$

Because it measures interaction between two frequency components – it retains information about the phase of the Fourier transform

Integrated Bispectrum

Phase of Fourier transform is a nonlinear function of frequency, extracted by biphas

$$\psi(\omega_1, \omega_2) = \phi(\omega_1) + \phi(\omega_2) - \phi(\omega_1 + \omega_2)$$

↑
Biphase

↑ ↑
Phase of Fourier spectrum

Need to find a feature identifier from bispectrum that can distinguish object shapes within the images - invariant under translation, scaling, rotation, and amplification

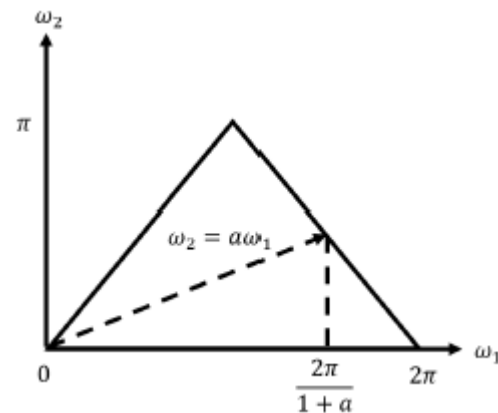
Define the integrated bi-spectrum:

$$I(a) = I_r(a) + i I_i(a) = \int_{\omega_1=0}^{2\pi/(1+a)} B(\omega_1, a\omega_1) d\omega_1 \quad \text{for } 0 < a < 1$$

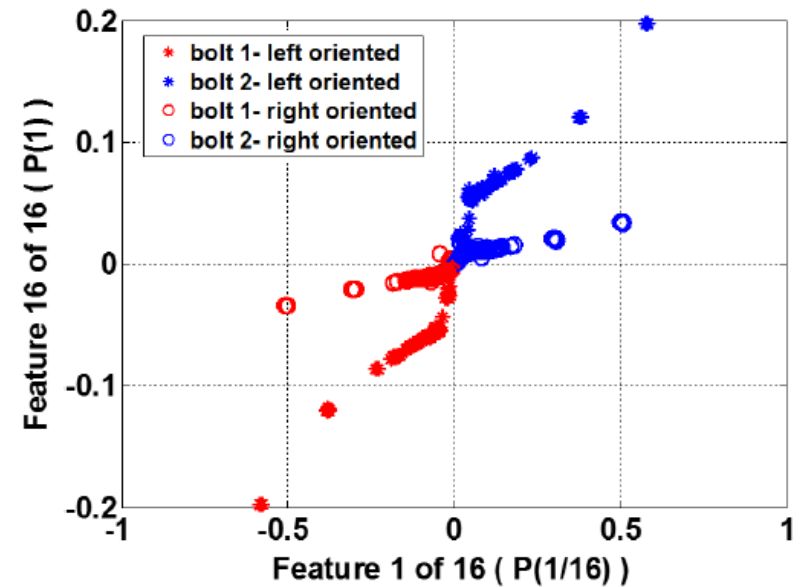
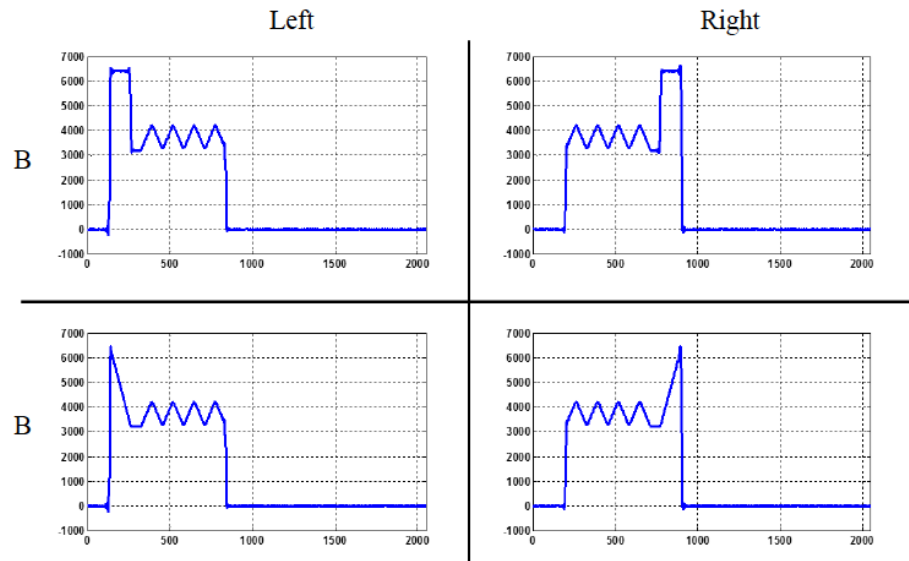
Define phase of integrated bi-spectrum

$$P(a) = \arctan\left(\frac{I_i(a)}{I_r(a)}\right)$$

Identifier $P(a)$ is invariant to translation, scaling, rotation and amplification
(Nikias and Raghuveer, 1987 and others)



Detecting Shapes



Bispectrum from sparse data

So far: Higher-order spectra can provide some interesting insights into spatial characteristics of reservoir models

However, FFT requires an exhaustive dataset or image (such as a training image)

What about when only scattered data is available?

Consider the inverse transform:

$$y_{N \times 1} = A_{N \times N} x_{N \times 1}$$

x is the Fourier transform coefficient, A is the inverse Fourier transform matrix, and y is the actual image. Because image exhibits spatial correlation, its Fourier transform is sparse.

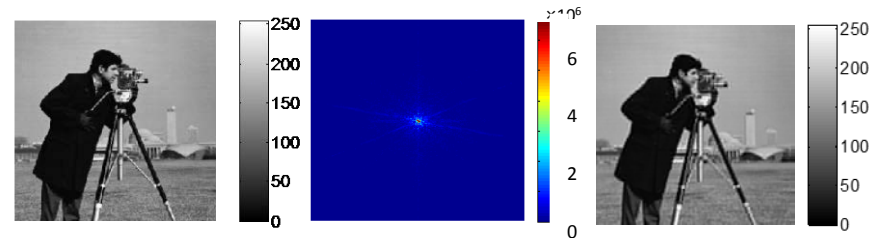


Image reconstructed using only highest 15% of Fourier coefficients

Compressed Sensing

In reservoir modeling sparse observations $y_{K \times 1}$ can be reconstituted as:

$$y_{K \times 1} = A_{K \times N} x_{N \times 1}$$

Problem is ill-posed and has multiple solutions. For a unique solution impose sparsity constraint on x

$$\min_{x \in \mathbb{R}^N} \|x\|_1 \quad \text{subject to} \quad y_{K \times 1} = A_{K \times N} x_{N \times 1}$$

Spatial correlation of geologic models, leads to x having a sparse representation  minimization problem

used for geostatistical simulation (Jafarpour, Goyal, McLaughlin, & Freeman, 2009) assuming a DCT

 unique solution – no uncertainty quantification

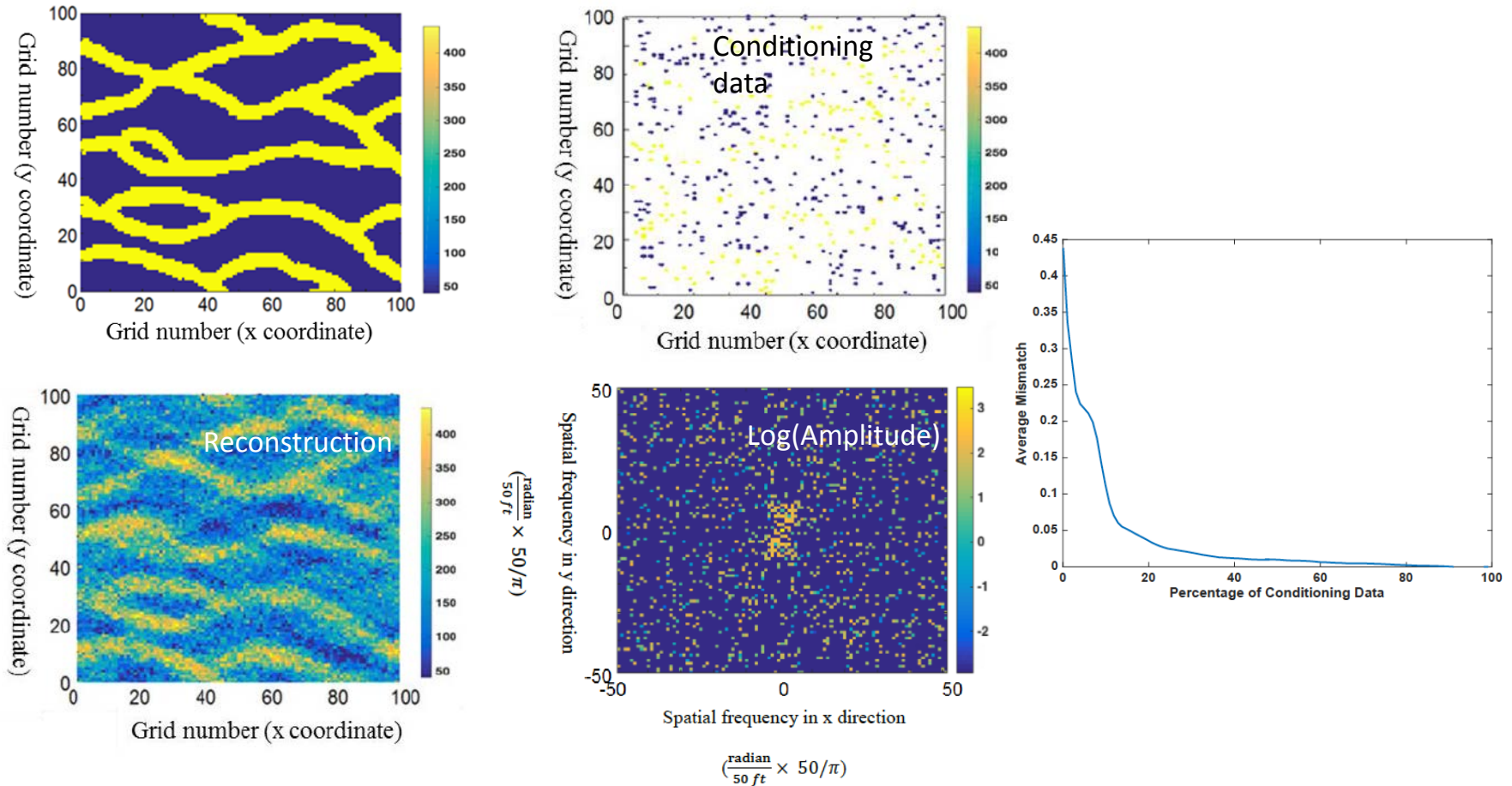
Sparse reconstruction

Estimate Fourier transform coefficients from scattered data
solving a regularization problem

$$\min_{x \in \mathbb{R}^N} \|y - Ax\|_p + \lambda \|x\|_q \quad \text{least mixed norm (LMN) [p=2, q=1]}$$

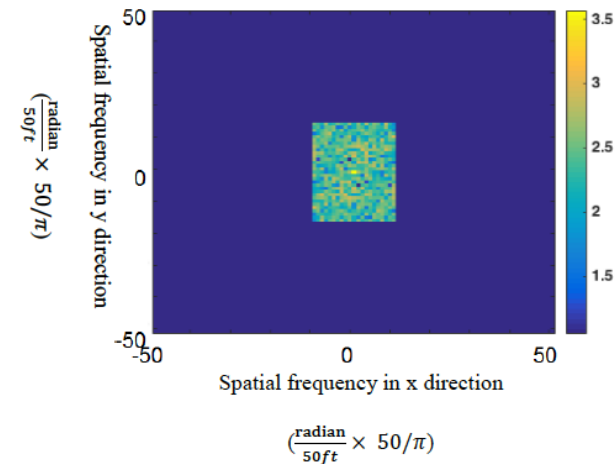
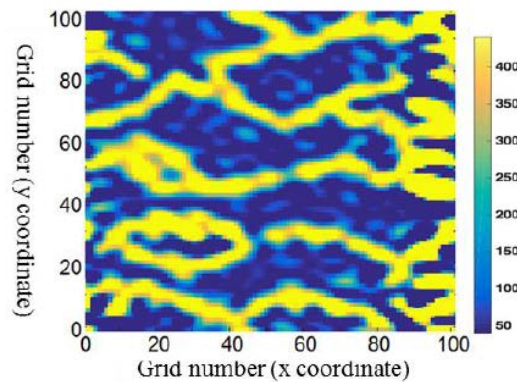
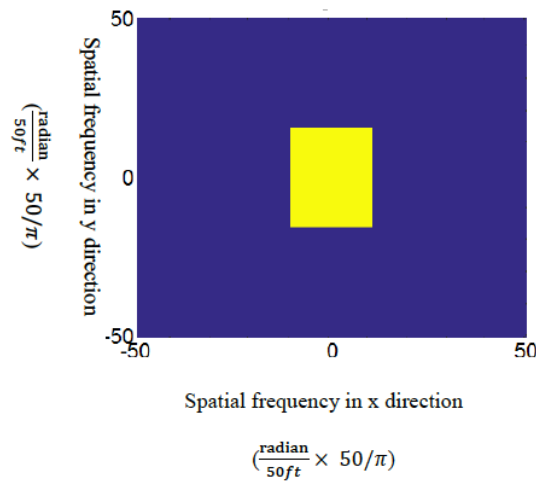
Optimum lambda established using jack-knife

Sparse reconstruction – some results



An improvement

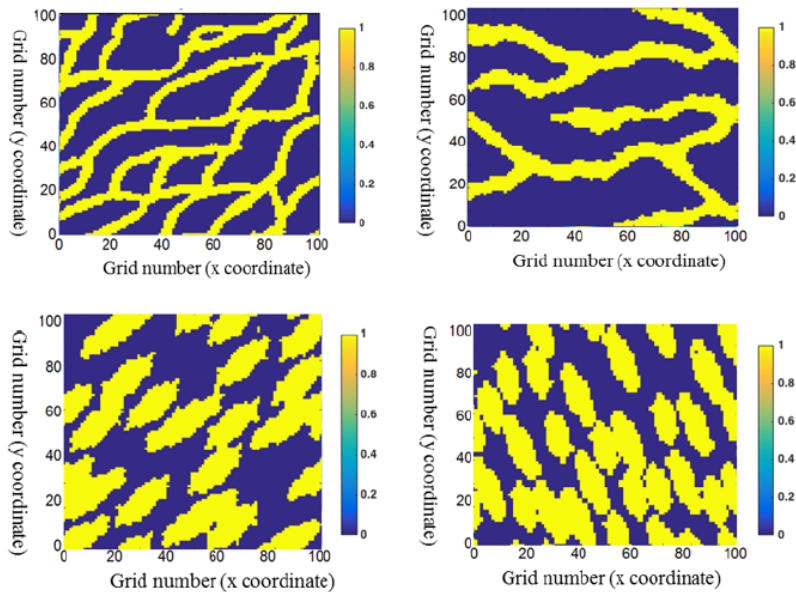
One idea to improve quality of reconstruction - limit the search space to low frequency coefficients



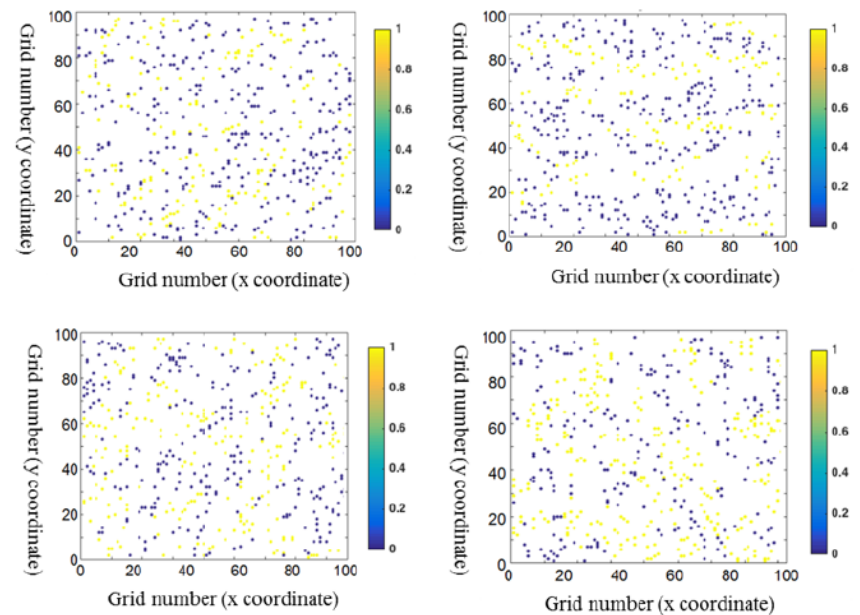
How to automatically select the search space? Apply a jackknife method.
Randomly select 10% of the conditioning data (validation data) and use the remaining data for reconstruction

Can we detect shapes using sparse data?

Reference Models

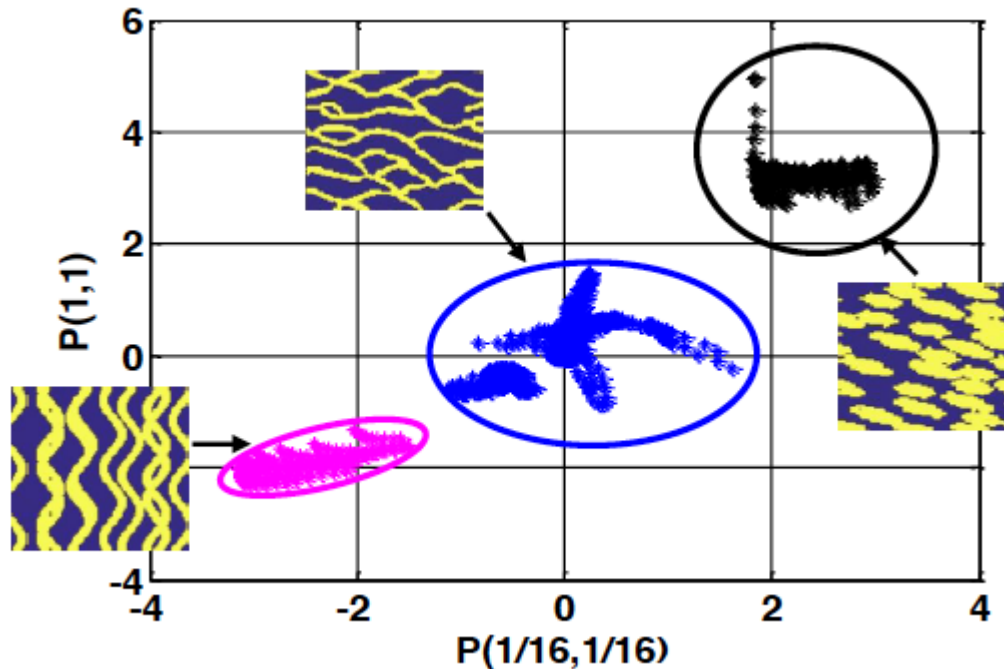


Sparse Conditioning Data



Can we detect shapes using sparse data?

Our goal is to use the integrated bi-spectrum computed using the Fourier transform from sparse data to classify the models reflecting different features~



Conclusions

- Performing inference and modeling of connectivity statistics in the Fourier space provides an avenue for modeling geological realism in a computationally efficient manner
- Computation of higher order moments, cumulant and polyspectra using sparse data appears feasible
- Detection of size, orientation and shapes of features consistent with available conditioning data is an interesting aspect of this research → training image selection
- While a broad idea about reservoir structures is possible using the power spectrum, the location of objects and details of the shapes are only possible by identifying phase → bispectrum, higher-order spectra