



# Understanding exchanges across stratified/convective zones interfaces

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#### Motivations...

<u>Generic situation:</u> a turbulent convective fluid layer stands above or below a stably stratified one, with a sharp but deformable interface.

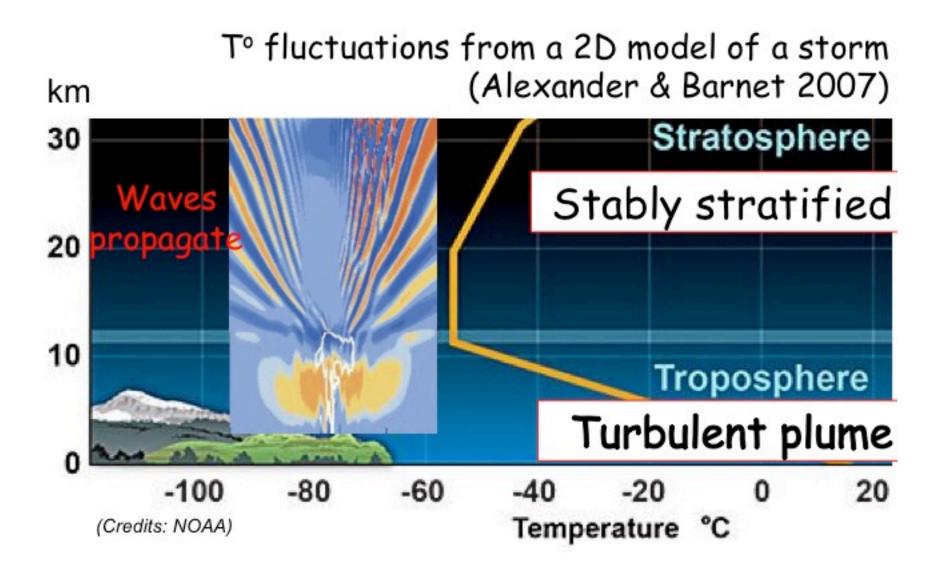
e.g. atmospheres, stars, planetary cores...

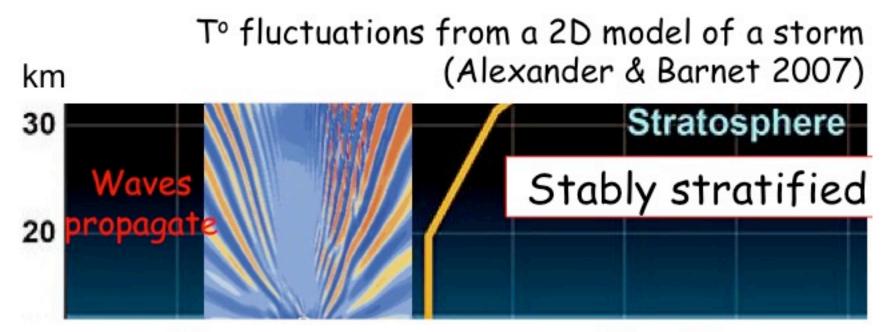
Turbulent plume



Storm







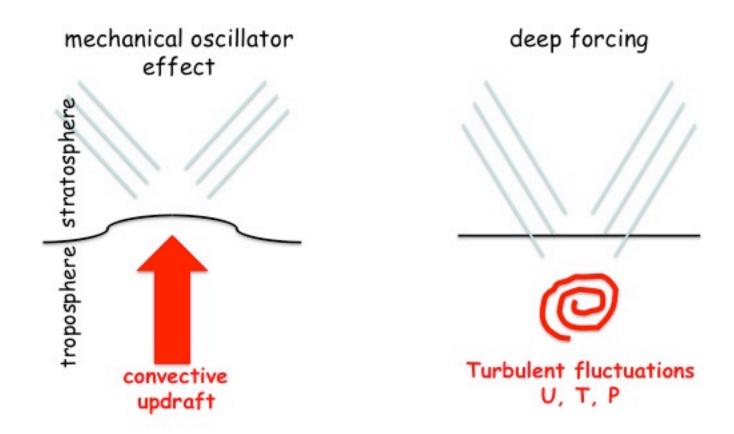
How are waves excited?

How and what do they propagate?

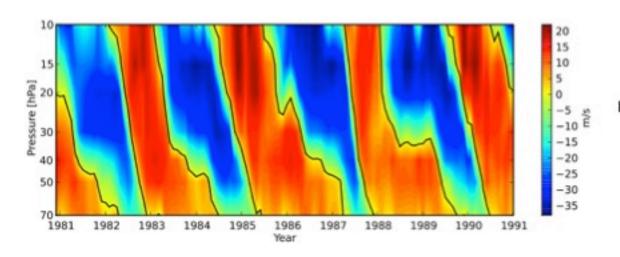
What are their consequences?

Gravity wave generation by convection

(e.g. Ansong & Sutherland 2010)



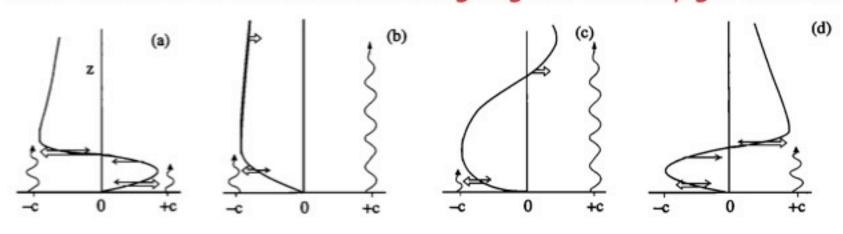
- Gravity waves affect the global energy budget
  - have to be included in general circulation models for accurate predictions of global weather patterns
  - > coarse grids with long time steps => parameterization
- Gravity waves carry momentum:
  - Breaking and mixing
  - Non-linear interactions generate zonal flows e.g. quasi-biennial oscillation (e.g. Plumb 1977)



monthly-mean zonalmean equatorial zonal wind in m/s between about 20 and 35 km

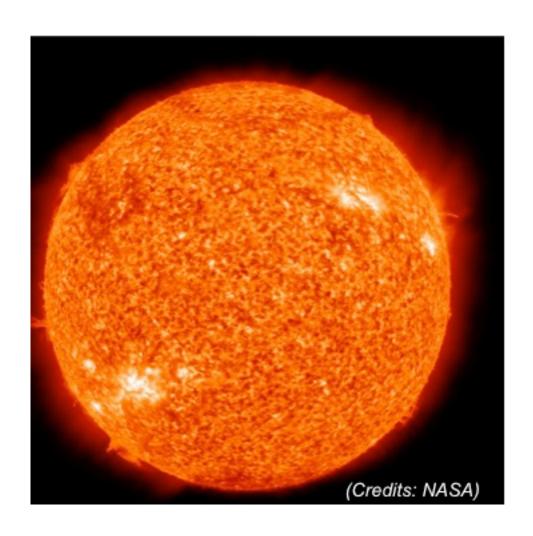
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anti-diffusive effect, accentuating angular velocity gradients...

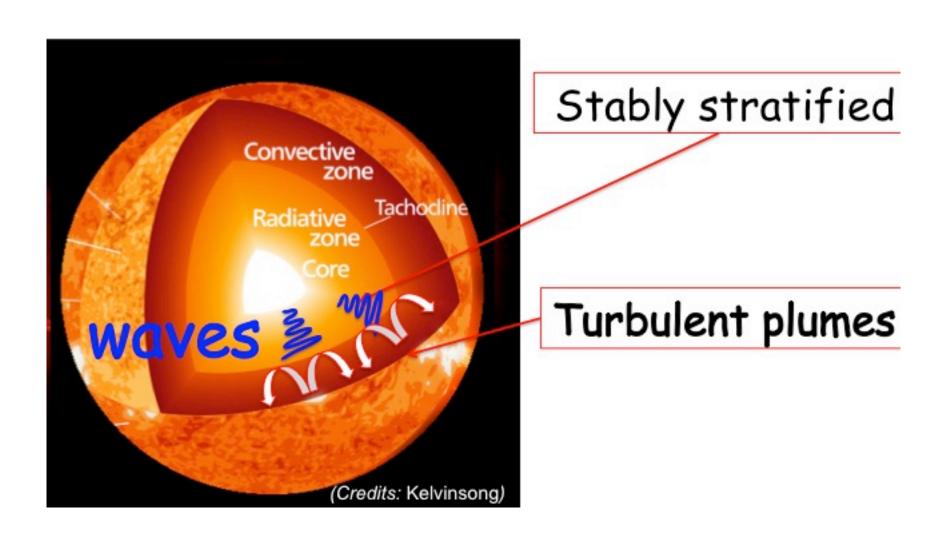


Stars...

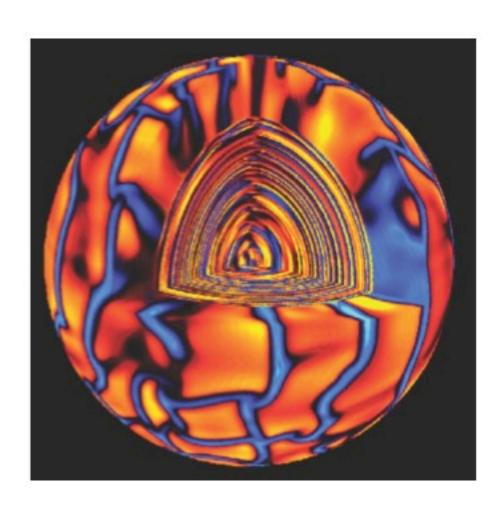
Same questions in stellar interiors...



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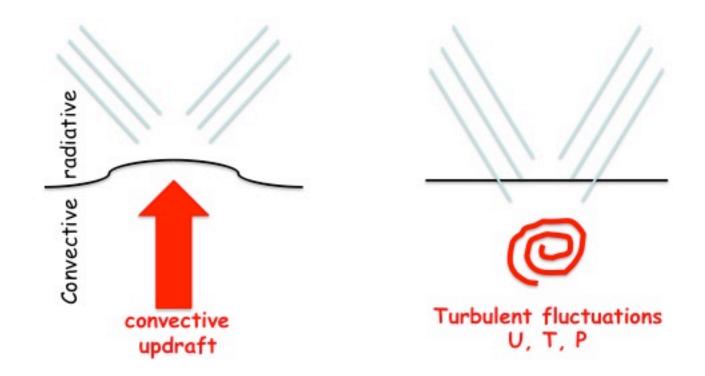
# Same questions in stellar interiors...



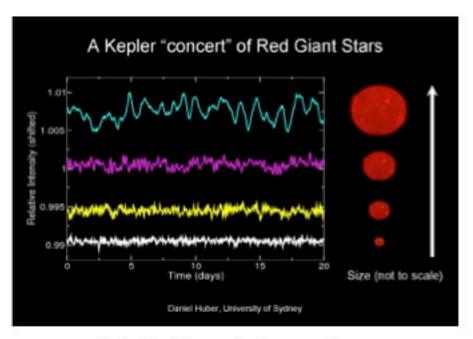
3D numerical simulations of the Sun by Alvan et al. (2012): time snapshot of the radial velocity.

Gravity waves excited

- · at the interface by overshooting plumes
- within the convective layer by Reynolds stress and entropy fluctuations



Gravity waves = important diagnostic tool of stellar structure in asteroseismology



Oscillations at the surface

Changes in the propagation

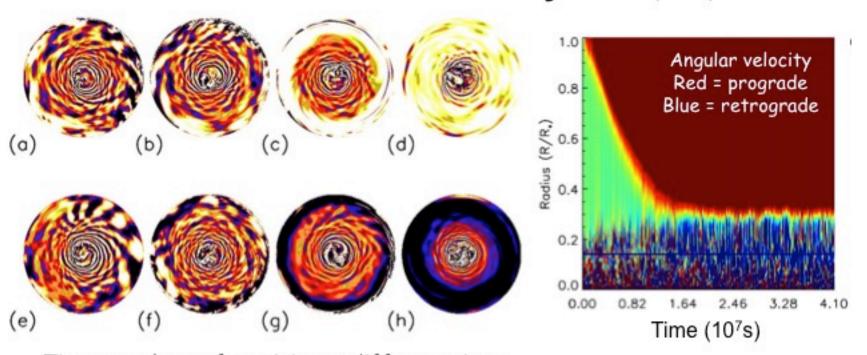
Internal structure

Gravity waves = important diagnostic tool of stellar structure in asteroseismology

Gravity waves transport energy and momentum:

- mixing, enhancing diffusion (e.g. observed lower Li abundance in F-stars, Charbonnel & Talon 2005)
- increased light flux and mass loss at the surface of massive starts (Quataert & Shiode 2012)
- selective damping because of symmetry breaking by rotation
  - => angular momentum deposit and modification of rotation profile.

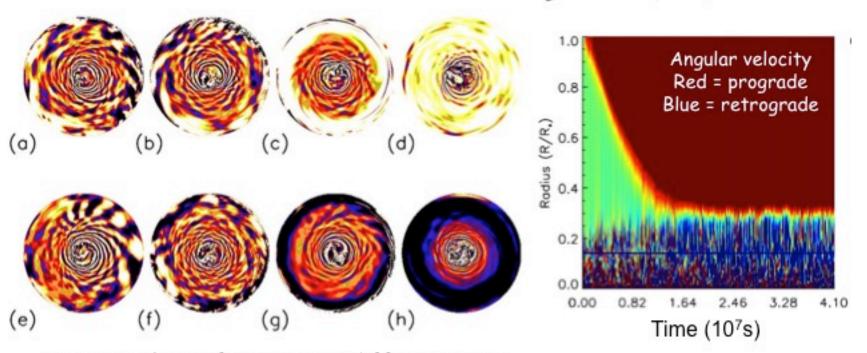
#### Axisymmetric numerical simulations by Rogers et al. (2012)



Time snapshots of vorticity at different times White = positive vorticity, black = negative vorticity

IGWs generally have an anti-diffusive effect, accentuating angular velocity gradients...

#### Axisymmetric numerical simulations by Rogers et al. (2012)



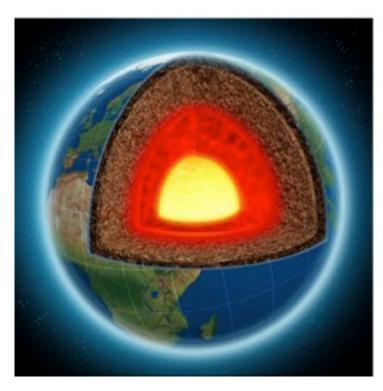
Time snapshots of vorticity at different times White = positive vorticity, black = negative vorticity

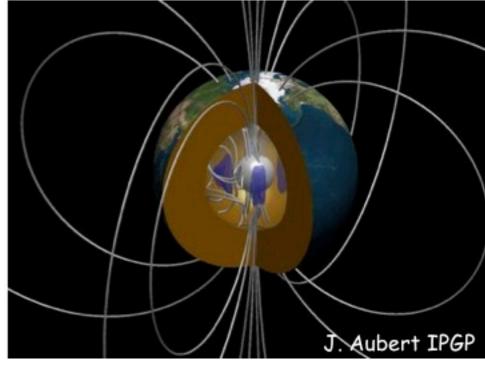
"Internal" mechanism for explaining the observed misalignment between extrasolar planets and their hot host stars

#### Earth's core...

#### Standard model:

- convective motions in the outer core
   mixing => adiabatic To and well-mixed composition.
- · growth of (nearly) pure iron inner core by crystallization
- energy budget = age of the inner core.





#### Earth's core...

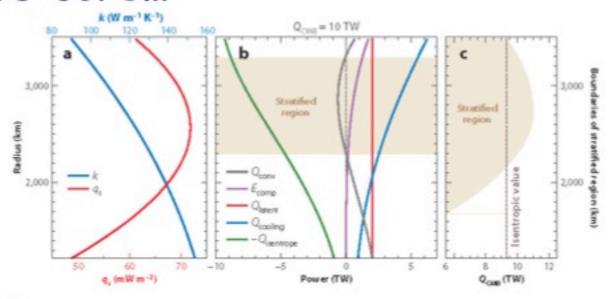
#### But

- ill-constrained composition of the core
  - ⇒ Density jump and temperature at ICB?
  - ⇒ Radiogenic heating?
- CMB heat flux over time? Present To (3800-4200K)?
- Extremely tight energy budget to drive convection: density difference  $\delta\rho/\rho\sim10^{-9}$   $\delta T_{lateral}\sim10^{-4}$  at CMB (Hirose et al 2013)
- Physical constants? (e.g. thermal conductivity)

#### possibility of stably stratified regions

- at the bottom and/or top of the outer core
- at intermediate depth...

#### Earth's core...



#### Figure 8

(a) Profiles of thermal conductivity (k) and heat flux along the isentrope (g<sub>t</sub>) in the core according to Gomi et al. (2011). (b) Convective heat flux (Q<sub>com</sub>, gray) is computed using an energy balance between the inner core boundary and each value of radius, for a core-mande boundary (CMB) heat flow of Q<sub>CMB</sub> = 10 TW. The balance writes as Q<sub>com</sub> + Q<sub>sectrops</sub> = E<sub>comp</sub> + Q<sub>latent</sub> + Q<sub>cosing</sub>, where E<sub>comp</sub> is the compositional energy due to light element transport in the chemical potential gradient, Q<sub>latent</sub> is the latent heat of inner core freezing, and Q<sub>cosing</sub> is the secular cooling of the shell. The region where the convective heat flow is negative (i.e., downward) tends to become stratified. The extent of this region is represented as a function of Q<sub>CMB</sub> in panel c.

Hirose et al. 2013

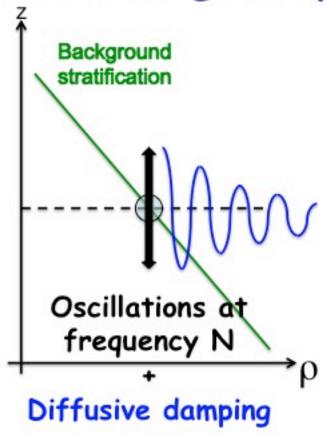
# Consequences on the geomagnetism?

# Open questions...

Turbulent convective patterns ?

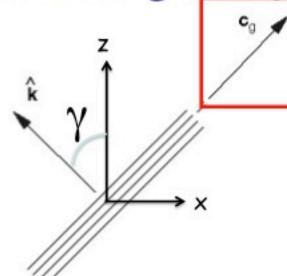
- · character of the wave field?
- amount of energy/momentum carried away?
- possible retro-action on the turbulence?
- Global integrated model & non-linear couplings
- Turbulence, stratification and waves
- Length and time scales spanning many orders of magnitude
  - > Numerics = very challenging
  - > Experimental study and analytical model

# Basics of gravity waves...



Fluid linearly stratified with density profile  $\rho = \rho_0 (1-N^2/g.z)$ , N being the buoyancy frequency

# Basics of gravity waves...



Generalisation: looking for plane wave solutions of the linearised Navier-Stokes equations

solutions = gravity waves 
$$(\mathbf{u}, \rho') = (\mathbf{u}_0, \rho_0') e^{i(\mathbf{k} \cdot \mathbf{r} - \omega t)}$$

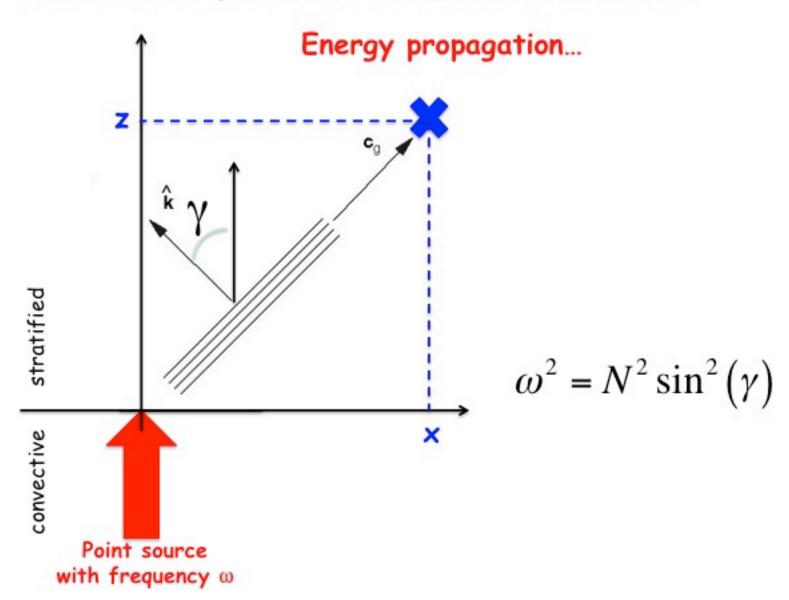
dispersion relation

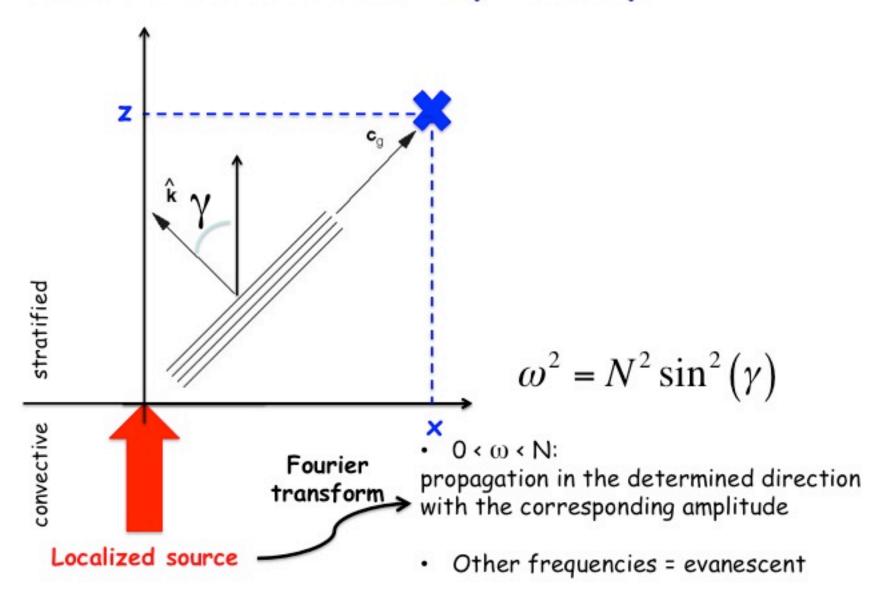
$$\omega^2 = N^2 \sin^2(\gamma)$$

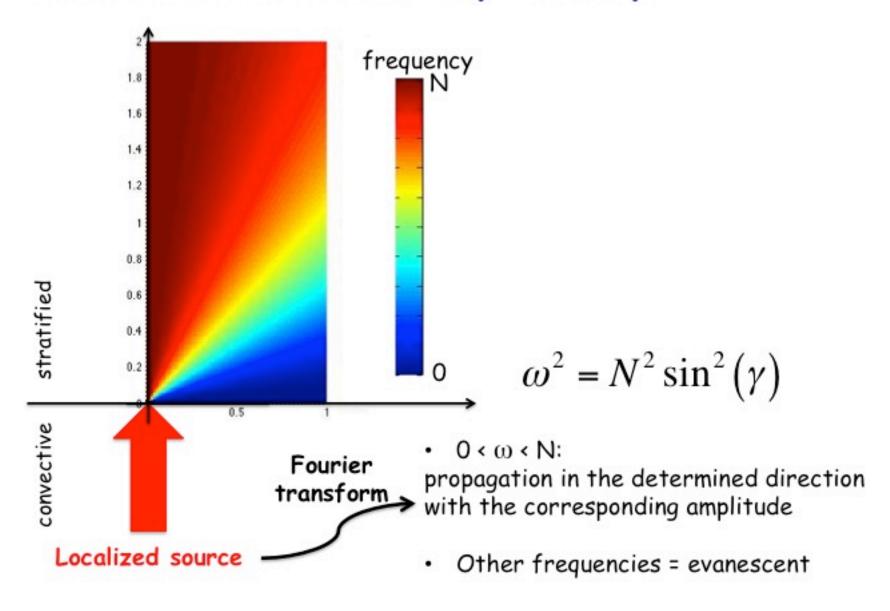
Group velocity
perpendicular to wavevector

$$\mathbf{k} \cdot \frac{\partial \omega}{\partial \mathbf{k}} = 0$$

## Emission from a localized source







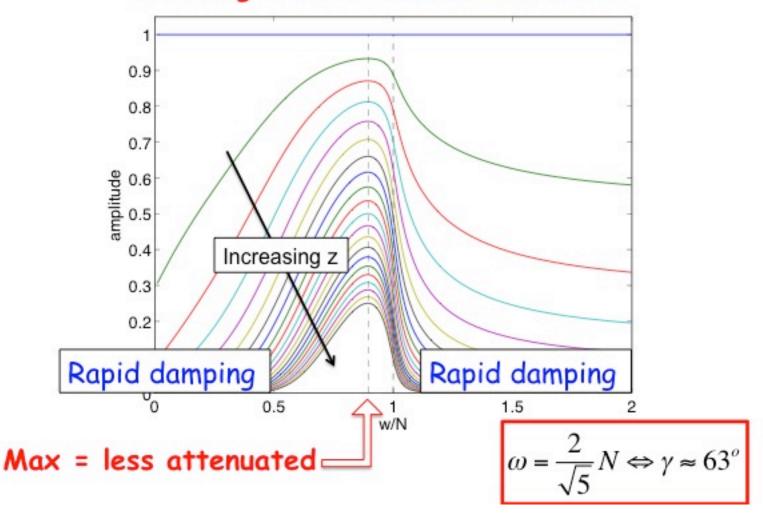
Including diffusive effects...

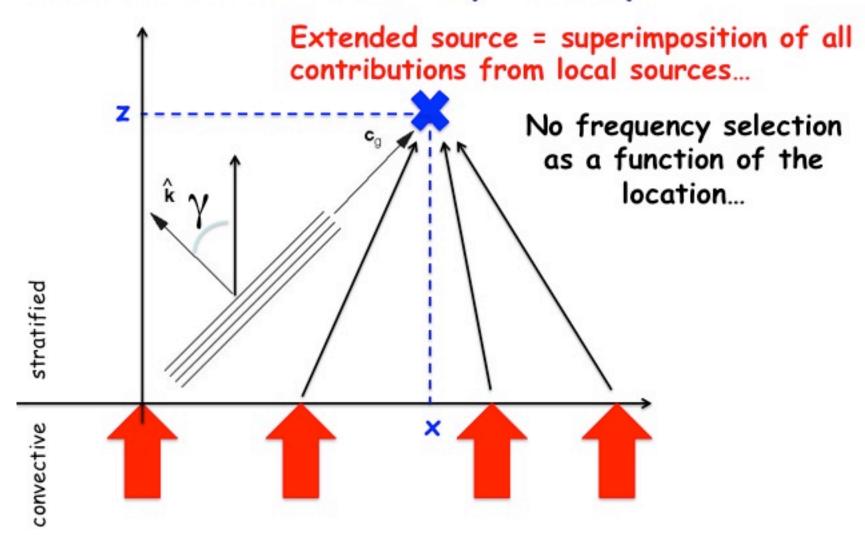
Diffusive dispersion relation: 
$$\left(i\omega + vk^2\right)^2 + N^2 \frac{\left(i\omega + vk^2\right)}{\left(i\omega + \kappa k^2\right)} k_x^2 = 0$$

Knowing the source signal (i.e.  $\omega$  and  $k_x$  at z=0), we solve for  $k_z$ 

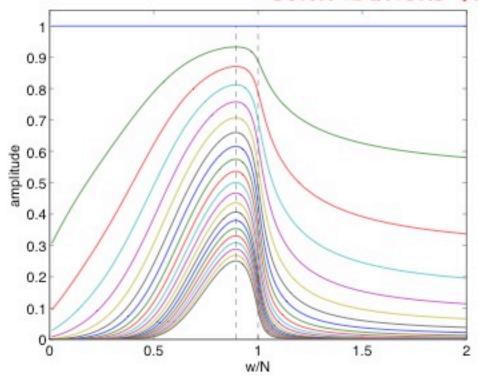
$$u=u_0e^{i(k.r-\omega t)}e^{-Att(v,\kappa,\gamma)z}$$

Wave amplitude at a given depth z as a function of  $\omega$  starting from a uniform excitation





Extended source = superimposition of all contributions from local sources...

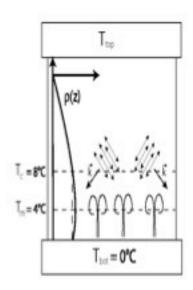


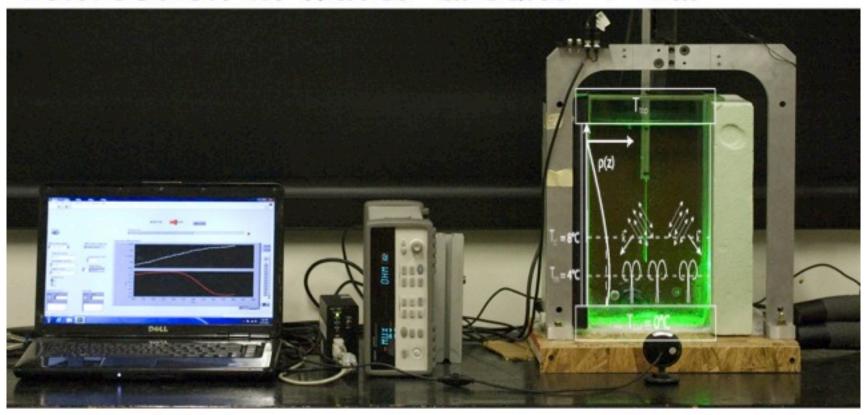
No frequency selection as a function of the location...

But still attenuation with depth and selective damping

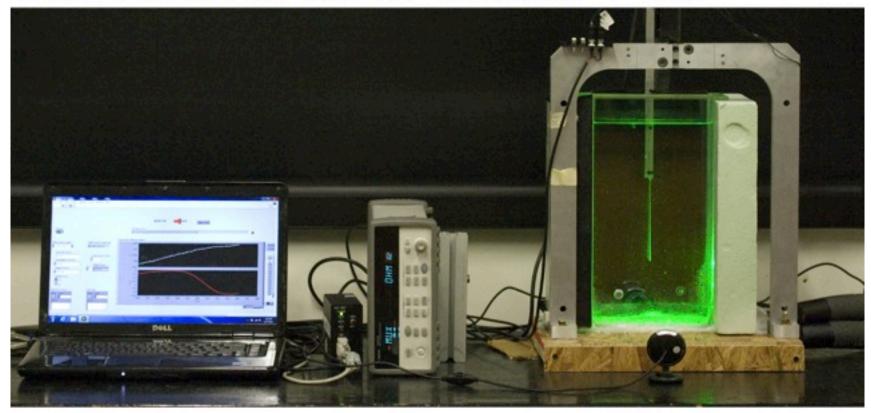
Validation in a simple, "self-organized", experimental system?

Water = maximum density at 4°C (Townsend 1964)

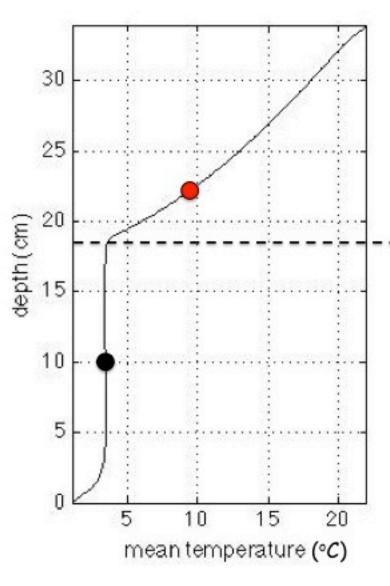




Dimensions:  $20 \times 4 \times 35 \text{ cm}^3$ 



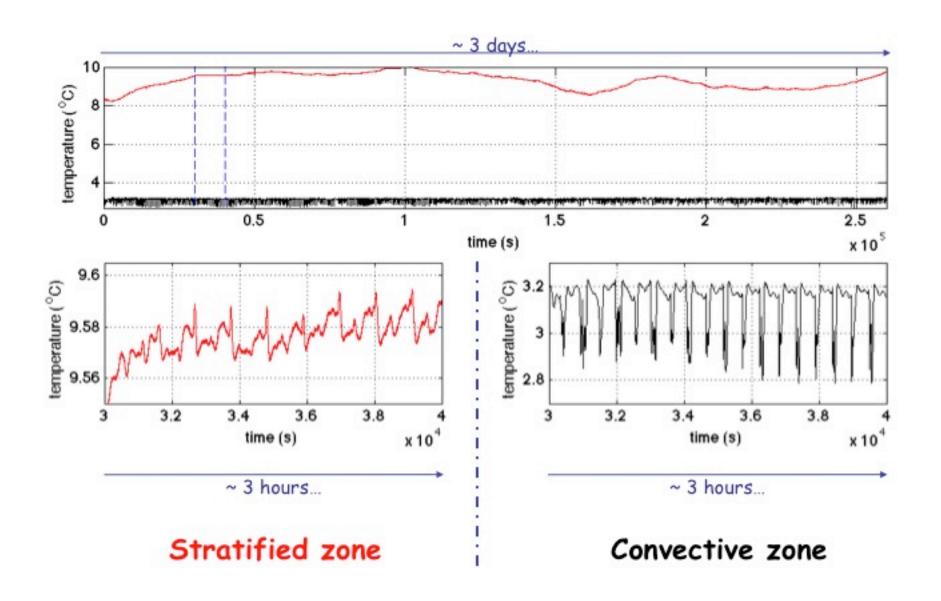
- Local T° measurements = high precision thermistors
- · PIV measurements in the convective and stratified zones

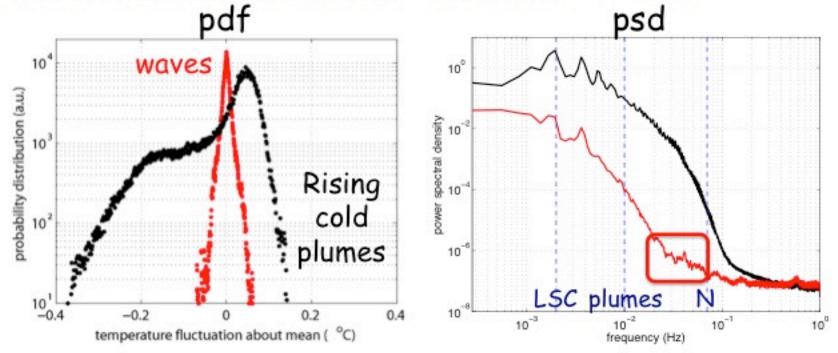


Stratified zone = heat transfer by conduction

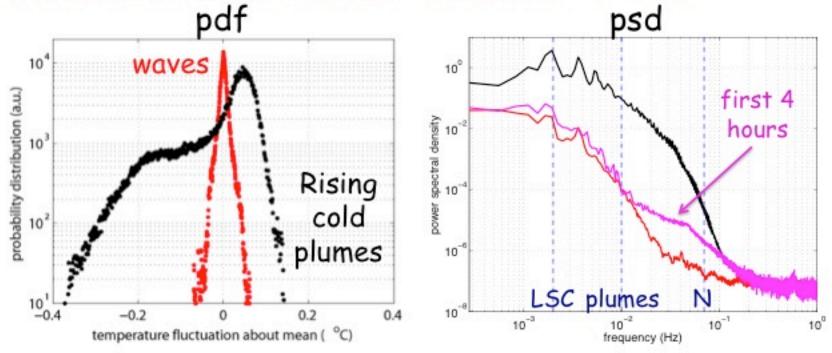
Interface (location function of imposed T° and heat losses)

Convective zone = heat transfer by convection Mean To ~3.4°C





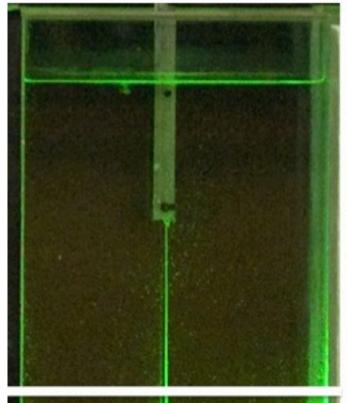
- Local excitation by random cold plumes
  - Gravity waves propagation
- Most energy in low frequency waves  $\leftarrow$  convective excitation
- No wave for f>N
- Less damping for high frequency waves



- Local excitation by random cold plumes
  - Gravity waves propagation
- Most energy in low frequency waves 

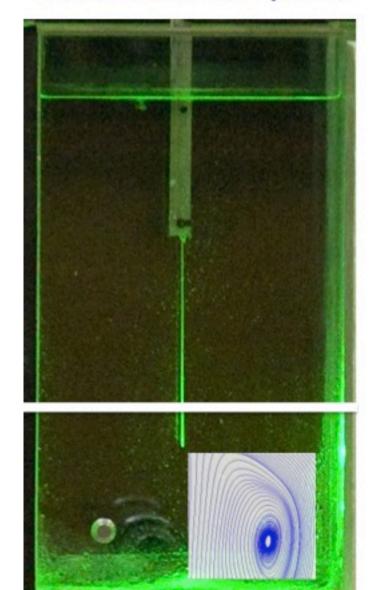
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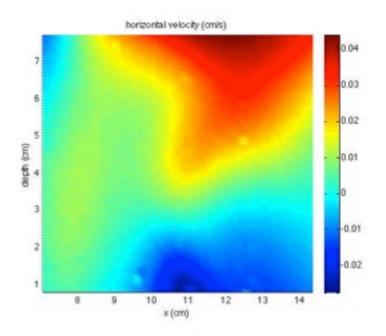
# PIV measurements

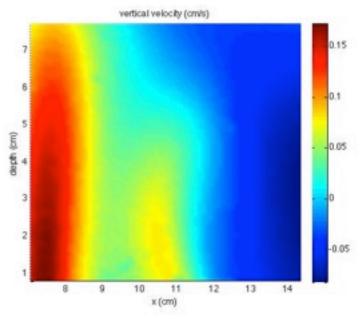


Interface location

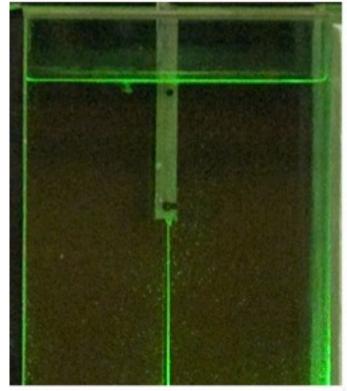
# Convective flow

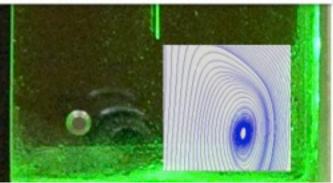




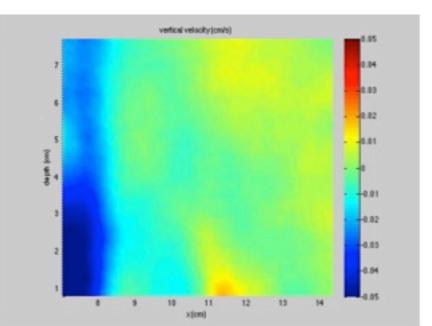


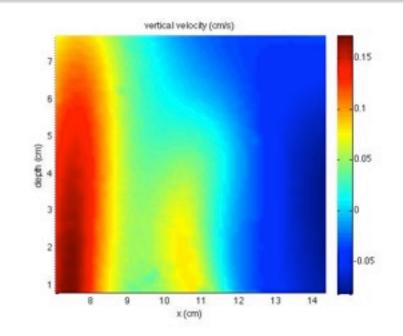
# Convective flow



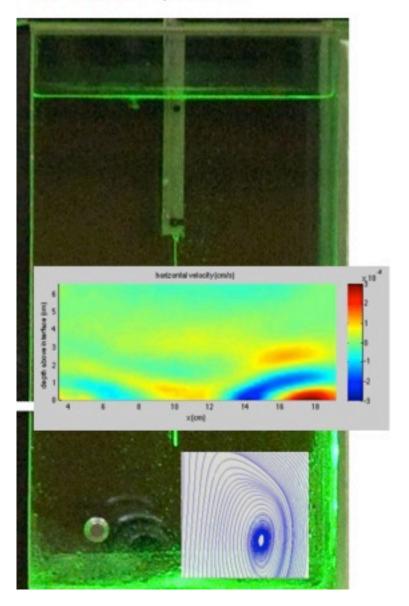


Accelerated x10 (real duration 10 min)

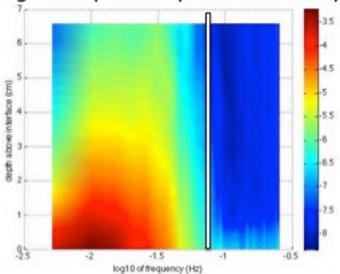




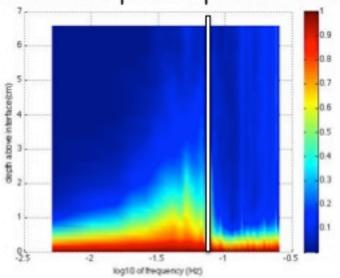
# Wave field



#### Log10 of power spectral density

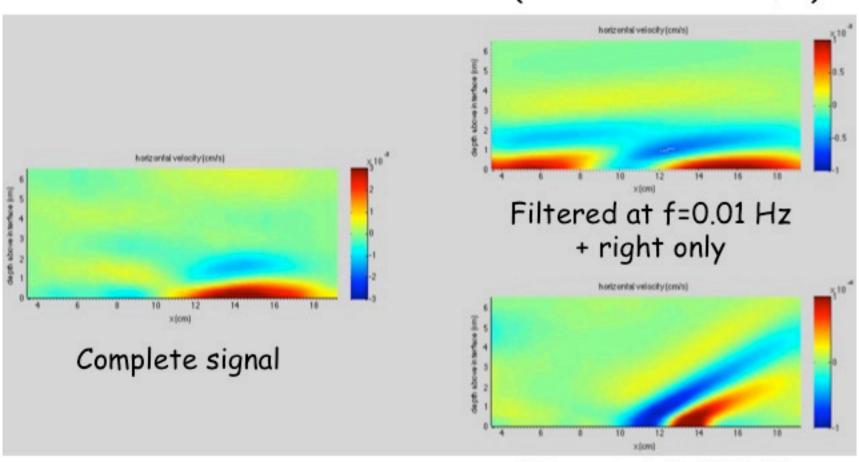


#### Renormalized power spectral density



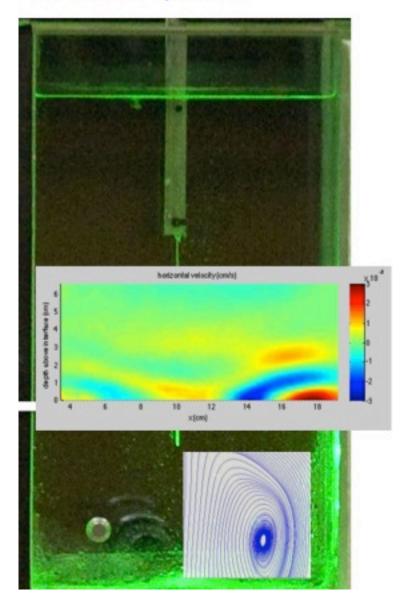
### Wave field

# Accelerated x20 (real duration = 13 min)

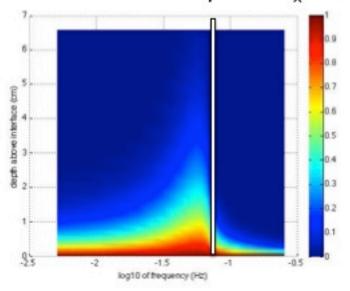


Filtered at f=0.06 Hz + right only

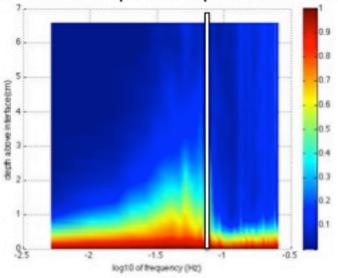
# Wave field

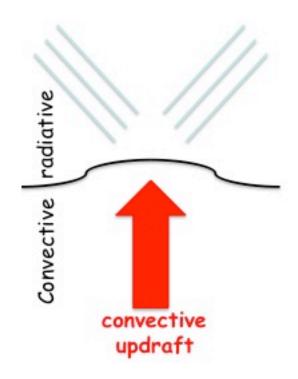


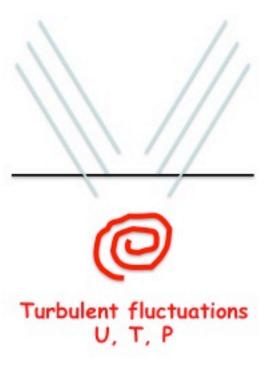
#### Linear diffusive theory with $k_{\rm x} \sim {\rm f}^{0.6}$



#### Renormalized power spectral density



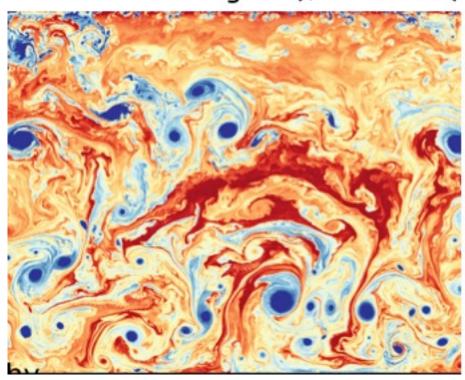




Numerical simulation...

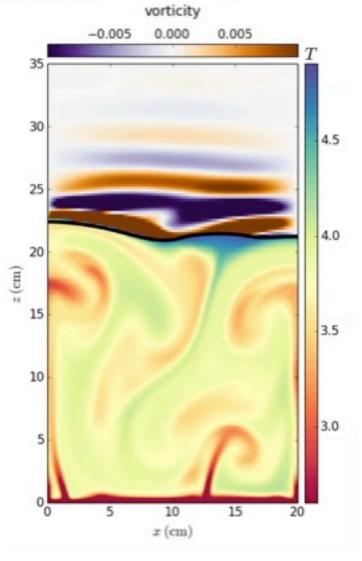
### Dedalus

D. Lecoanet (Berkeley), K. Burns (MIT), J. Oishi (AMNH/ SUNY Farmingdale), B. Brown (Colorado), G. Vasil (Sydney)



Pseudo-spectral
Open-source
Python
Very flexible equations

Numerical simulation... t= 11260.455 (s)



Simulation of the simulation: use full simulation data as inputs for simplified model simulations

# Deep Forcing (Lighthill)

Split full solution into linear convective and linear wave modes:

$$\boldsymbol{u} = \boldsymbol{u}_c + \partial_t \boldsymbol{\xi}$$

Project onto linear wave mode:

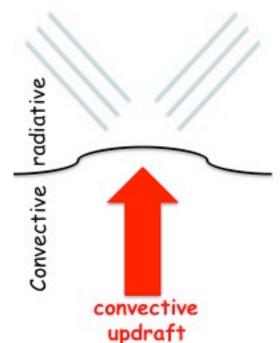
$$\nabla^2 \partial_t^2 \xi_z - N^2(z) \nabla_\perp^2 \xi_z = S$$



Turbulent fluctuations U, T, P

from water simulation

Simulation of the simulation: use full simulation data as inputs for simplified model simulations



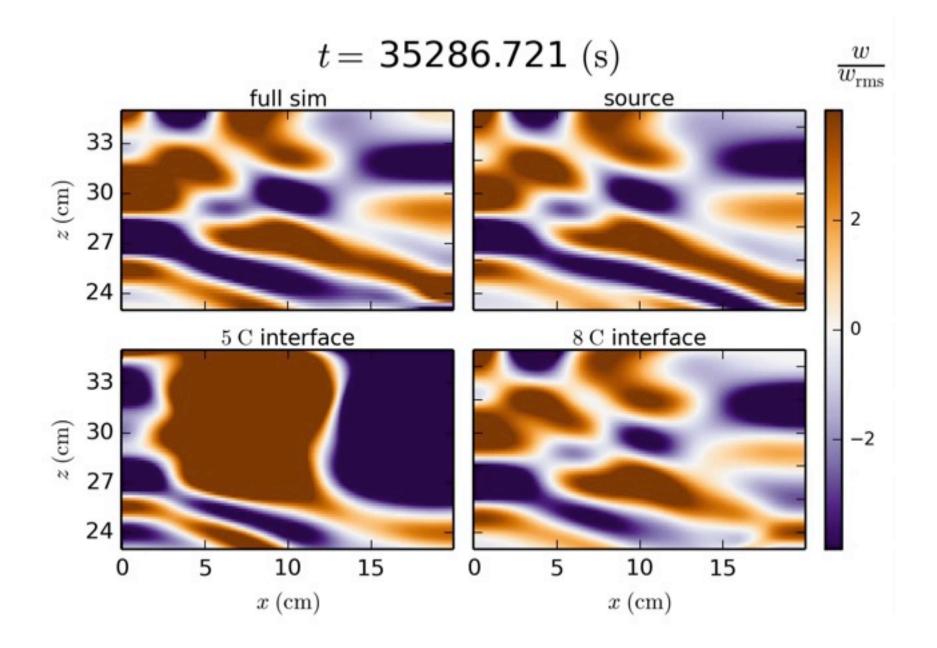
### Interface fluctuations

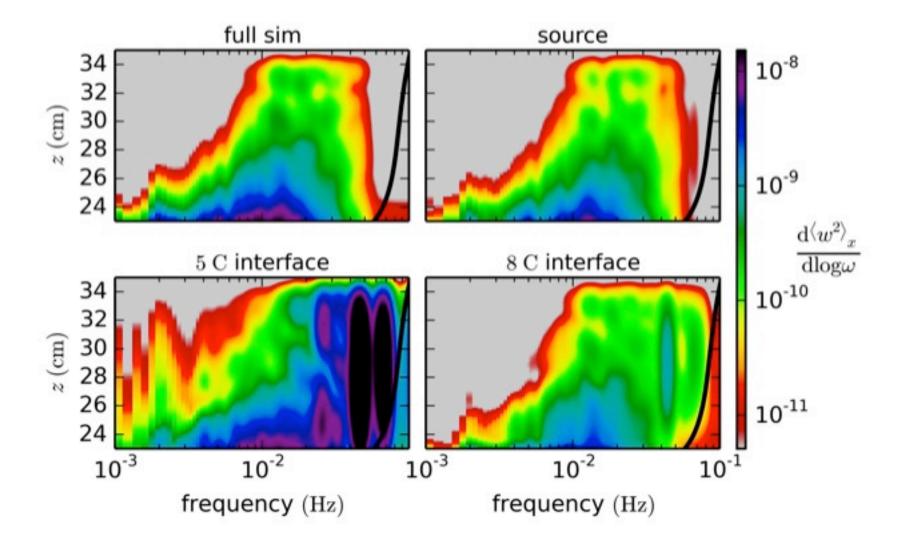
No source term, but force boundaries

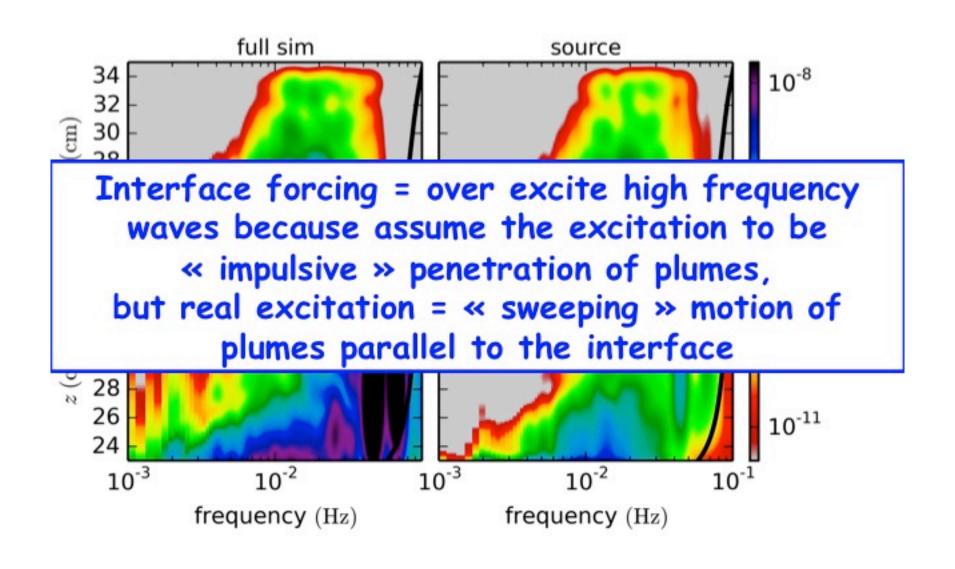
$$\nabla^2 \partial_t^2 \xi_z - N^2(z) \nabla_\perp^2 \xi_z = 0$$

Boundary condition:

$$\xi_z(x, z_{\rm int}) = z_{\rm int}(x) - \bar{z}_{\rm int}$$





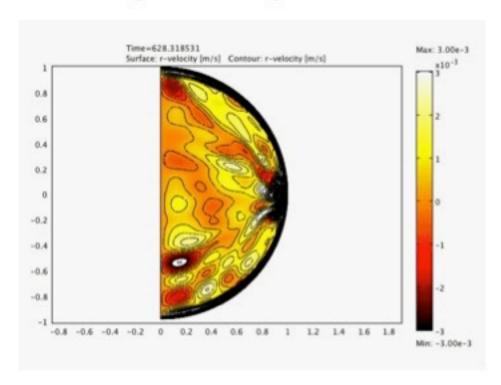


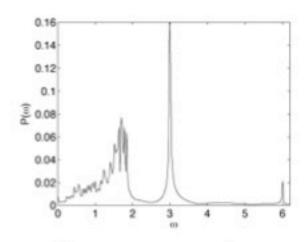
#### Conclusions

- A stratified region above/below a turbulent one is not motionless, but carries part of the energy
- Required = statistics of the convective source
- then wave amplitude and frequency selection depend on diffusive processes
- Main excitation mechanism = Reynolds stress... Can use this
  to analytically estimate result for more turbulent cases
  (e.g., stars, see Lecoanet & Quataert 2013)
- Coming studies:
  - > generation of a mean flow? Generalization of the QBO...
  - effect of global rotation...

### Conclusions

Generalization to other turbulence sources...
 e.g. boundary turbulence in librating planets





⇒ frequency selection in inertial waves (Sauret et al. 2013)