NON-LINEARITIES IN GEODYNAMO MODELS AND THEIR CONNECTION TO THE WEAK-FIELD AND STRONG-FIELD BRANCHES

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## The Earth's internal structure



$$\left(\frac{\partial \mathbf{u}}{\partial t} + (\mathbf{u} \cdot \nabla)\mathbf{u}\right) = -\nabla\pi + \nu \,\Delta \mathbf{u} - 2\Omega(\mathbf{e}_z \wedge \mathbf{u})$$
$$-\alpha T\mathbf{g}$$

 $\begin{aligned} \boldsymbol{\nabla} \cdot \mathbf{u} &= 0 \,, \\ \frac{\partial T}{\partial t} + \mathbf{u} \cdot \boldsymbol{\nabla} T &= \kappa \, \Delta T \,, \end{aligned}$ 



# $\begin{aligned} \frac{\partial \mathbf{B}}{\partial t} &= -\boldsymbol{\nabla} \times \mathbf{E} \,, \qquad \boldsymbol{\nabla} \cdot \mathbf{B} = 0 \,, \qquad \boldsymbol{\nabla} \wedge \mathbf{B} = \mu_0 \, \mathbf{j} \,, \\ \mathbf{j} &= \sigma \mathbf{E}' = \sigma \left( \mathbf{E} + \mathbf{u} \wedge \mathbf{B} \right) \,, \end{aligned}$



$$\left(\frac{\partial \mathbf{u}}{\partial t} + (\mathbf{u} \cdot \nabla)\mathbf{u}\right) = -\nabla\pi + \nu \,\Delta \mathbf{u} - 2\Omega(\mathbf{e}_z \wedge \mathbf{u})$$
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$$-\alpha T \mathbf{g}$$

$$\begin{aligned} \nabla \cdot \mathbf{u} &= 0, \\ \frac{\partial T}{\partial t} + \mathbf{u} \cdot \nabla T &= \kappa \, \Delta T, \\ \frac{\partial \mathbf{B}}{\partial t} &= \nabla \wedge (\mathbf{u} \wedge \mathbf{B}) + \eta \, \Delta \mathbf{B}, \end{aligned}$$

 $\boldsymbol{\nabla}\cdot\mathbf{B}=0.$ 

$$\begin{cases} \left(\frac{\partial \mathbf{u}}{\partial t} + (\mathbf{u} \cdot \nabla) \mathbf{u}\right) = -\nabla \pi + \nu \, \Delta \mathbf{u} - 2\Omega(\mathbf{e}_z \wedge \mathbf{u}) \\ - \alpha \, T \mathbf{g} + \frac{1}{\mu_0 \rho_0} \left(\nabla \wedge \mathbf{B}\right) \wedge \mathbf{B} , \\ \nabla \cdot \mathbf{u} = 0 , \\ \frac{\partial T}{\partial t} + \mathbf{u} \cdot \nabla T = \kappa \, \Delta T , \\ \frac{\partial \mathbf{B}}{\partial t} = \nabla \wedge (\mathbf{u} \wedge \mathbf{B}) + \eta \, \Delta \mathbf{B} , \\ \nabla \cdot \mathbf{B} = 0 . \end{cases}$$

$$\begin{cases} E_{\eta} \left( \frac{\partial \mathbf{u}}{\partial t} + (\mathbf{u} \cdot \nabla) \mathbf{u} \right) = -\nabla \pi + E \Delta \mathbf{u} - \mathbf{e}_{z} \wedge \mathbf{u} \\ - Ra T \mathbf{g} + (\nabla \wedge \mathbf{B}) \wedge \mathbf{B}, \\ \nabla \cdot \mathbf{u} = 0, \\ \frac{\partial T}{\partial t} + \mathbf{u} \cdot \nabla T = q \Delta T, \\ \frac{\partial \mathbf{B}}{\partial t} = \nabla \wedge (\mathbf{u} \wedge \mathbf{B}) + \Delta \mathbf{B}, \\ \nabla \cdot \mathbf{B} = 0. \end{cases}$$
$$E/Pm = E_{\eta} = \frac{\eta}{2\Omega \mathcal{L}^{2}}, \quad E = \frac{\nu}{2\Omega \mathcal{L}^{2}}, \quad q = \frac{\kappa}{\eta}.$$

$$2\Omega \mathcal{L}^2$$
  $2\Omega \mathcal{L}^2$ 

$$\begin{cases} E_{\eta} \left( \frac{\partial \mathbf{u}}{\partial t} + (\mathbf{u} \cdot \nabla) \mathbf{u} \right) = -\nabla \pi + \mathbf{E} \Delta \mathbf{u} - \mathbf{e}_{z} \wedge \mathbf{u} \\ - \mathbf{R} a T \mathbf{g} + (\nabla \wedge \mathbf{B}) \wedge \mathbf{B}, \\ \nabla \cdot \mathbf{u} = 0, \\ \frac{\partial T}{\partial t} + \mathbf{u} \cdot \nabla T = \mathbf{q} \Delta T, \\ \frac{\partial \mathbf{B}}{\partial t} = \nabla \wedge (\mathbf{u} \wedge \mathbf{B}) + \Delta \mathbf{B}, \\ \nabla \cdot \mathbf{B} = 0. \end{cases}$$

$$\frac{E/Pm = E_{\eta}}{\eta} = \frac{\eta}{2\Omega \mathcal{L}^2}, \quad E = \frac{\nu}{2\Omega \mathcal{L}^2}, \quad q = \frac{\pi}{\eta}.$$
$$\simeq 10^{-7} \qquad \simeq 10^{-15} \qquad \simeq 10^{-7}$$

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#### Weak and Strong field branches

Chandrasekhar, 1961 Eltayeb & Roberts, 1970 Eltayeb & Kumar, 1977 Roberts & Soward, 1978 Soward, 1979 Fearn, 1979 Fautrelle & Childress, 1982 Proctor & Weiss, 1982

## Weak and Strong field branches



## Weak and Strong field branches



(Roberts, GAFD, 1988)

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# Glatzmaier & Roberts, 1995



#### Roberts number q = Pm / Pr





• E<sub>ma</sub>/ \* No<sup>g</sup>dynamo

# Christensen-Aubert 2006



## Power based scaling laws for the magnetic field



Statistical energy balance between production and dissipation

$$P^{\star} = D_{\nu}^{\star} + D_{\eta}^{\star} \qquad f_{ohm} P^{\star} = D_{\eta}^{\star}$$
where  $D_{\eta}^{\star} = E_{\eta} \int_{V} (\nabla \times \mathbf{B}^{\star})^{2} dV^{\star}$ 
 $f_{ohm} P^{\star} = E_{\eta} \frac{\mathrm{Lo}^{2}}{\ell_{B}^{\star}^{2}}$ 

 $\ell_B^{\star}$ : magnetic dissipation length scale  $\ell_B^{\star 2} \equiv \frac{\int_V \mathbf{B}^{\star 2} \, \mathrm{d}V^{\star}}{\int_V (\nabla \times \mathbf{B}^{\star})^2 \, \mathrm{d}V^{\star}}$ 

Oruba & Dormy, GJI, 198, 828-847 (2014).

#### Statistical energy balance between production and dissipation

$$f_{\rm ohm} \mathbf{P}^{\star} = \mathbf{E}_{\eta} \frac{\mathrm{Lo}^2}{\ell_B^{\star 2}}$$

$$- Lo = f_{ohm}^{1/2} P^{\star 1/2} E_{\eta}^{-1/2} \ell_B^{\star}$$

#### Verified by any statistically steady dynamo

Oruba & Dormy, GJI, 198, 828-847 (2014).

#### Statistical energy balance between production and dissipation

The simplest approximation:  $\ell_B^{\star}$  constant  $\text{Lo} \sim f_{\text{ohm}}^{1/2} \operatorname{Ra}_Q^{\star}^{1/2} \operatorname{E}_{\eta}^{-1/2}$ 



#### Power based scaling laws for the magnetic field

• Energy balance +  $\ell_B^{\star}$  constant + more precise description of  $\ell_B^{\star}$ :



## Power based scaling laws for the magnetic field

• Energy balance +  $\ell_B^{\star}$  constant + more precise description of  $\ell_B^{\star}$ :

$$Lo \sim f_{ohm}^{1/2} Ra_Q^{\star 1/2} E_{\eta}^{-1/2}$$

$$Lo \sim f_{ohm}^{1/2} Ra_Q^{\star 0.32} Pm^{0.11}$$



## Power based magnetic field scalings

# TOO GENERAL:

applicable to any model **irrespectively of** the magnetic field generation mechanism

Oruba & Dormy, GJI, 198, 828-847 (2014).

#### Viscous length scale of the flow



Oruba & Dormy, GJI, 198, 828-847 (2014).

#### Magnetic field strength

Bifurcation from a laminar flow (Fauve & Pétrélis, 2007): Lorentz force  $\sim$  modified viscous force

with  $\ell_u^{\star} \sim \mathrm{E}^{1/3}$ 

$$\Lambda \sim (\operatorname{Rm} - \operatorname{Rm}_d) \operatorname{E}^{1/3} \quad \Lambda = \frac{\sigma B^2}{\rho \Omega}$$
: Elsasser

D?



#### Magnetic field strength

Bifurcation from a laminar flow (Fauve & Pétrélis, 2007): Lorentz force  $\sim$  modified viscous force

with  $\ell_u^{\star} \sim \mathrm{E}^{1/3}$ 

10

10

10

10

(Reduced database)

 $Rm - Rm_d$ 

AE-

$$\Lambda \sim (\operatorname{Rm} - \operatorname{Rm}_d) \operatorname{E}^{1/3} \quad \Lambda = \frac{\sigma B^2}{\rho \Omega}$$
: Elsasser

Application to the Earth's core?  $\operatorname{Rm} - \operatorname{Rm}_d \le 10^3$  $\longrightarrow \Lambda \le 10^{-2}$ 

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much smaller than  $\Lambda_{\text{Earth}} \simeq 1$ 

Oruba & Dormy, GJI, 198, 828-847 (2014).

## Intermediate result

## Dipolar Dynamos at moderate forcing are dominated by a balance between viscous and Coriolis forces

#### Roberts number q = Pm / Pr





• E<sub>ma</sub>/ \* No<sup>g</sup>dynamo



## Transition from dipolar to multipolar regime

Dominance of the Coriolis force in both regimes:

Transition controlled by the relative strength of inertial to viscous forces:



Soderlund et al 2012

#### Inertial forces versus Coriolis

$$\{\boldsymbol{\nabla} \times (\mathbf{u} \times \boldsymbol{\nabla} \times \mathbf{u})\} \sim \{\boldsymbol{\nabla} \times (\mathbf{e}_z \times \mathbf{u})\}$$
$$\frac{\mathbf{u}^2}{\ell \ell_u} \sim \frac{\mathbf{u}}{\ell_{//}}$$

#### **Typical length scales**

the kinematic dissipation length scale  $\ell_u$ :

the length scale  $\ell$ 

$$a_u^2 \equiv \frac{\langle \mathbf{u}^2 \rangle}{\langle (\nabla \times \mathbf{u})^2 \rangle}$$

(Oruba & Dormy 2014)

the parallel length scale :  $\ell_{//} \sim 1$  (quasi-geostrophy)

 $\frac{\operatorname{Ro}_{\ell} \sim \ell}{\operatorname{Oruba} \& \operatorname{Dormy}, \operatorname{GRL} \text{ in press.}}$ 

#### Inertial forces versus Coriolis

Test with the numerical database (U. Christensen):


#### Viscous forces versus Coriolis



#### Inertial versus viscous forces

# $\{ \boldsymbol{\nabla} \times (\mathbf{u} \times \boldsymbol{\nabla} \times \mathbf{u}) \} \sim \{ \mathbf{E} \, \boldsymbol{\nabla} \times \boldsymbol{\nabla} \times \boldsymbol{\nabla} \times \mathbf{u} \}$ $\frac{\mathbf{u}^2}{\ell \, \ell_u} \sim \frac{\mathbf{E} \, \mathbf{u}}{\ell_u^3}$ $\frac{\mathbf{u}}{\mathbf{E}} \ell_u^2 \sim \ell$

 $\operatorname{Re} \ell_u^2 \sim \ell$ 

Oruba & Dormy, GRL in press.

#### Inertial versus viscous forces

Test against the numerical database:



#### A dominant three forces balance at the transition



# $\operatorname{Ro} \mathrm{E}^{-1/3} \equiv \operatorname{Re} \mathrm{E}^{2/3} \sim \ell$

Oruba & Dormy, GRL in press.

#### Unified description of the transition



#### Bistability between both branches



Simitev & Busse, Bistability and hysteresis of dipolar dynamos generated by turbulent convection in rotating spherical shells, *Europhysics Letters (EPL)*, 85 (2009) 19001

#### Bistability between both branches



## Intermediate results

- Dipolar Dynamos at moderate forcing are dominated by a balance between viscous and Coriolis forces

 Loss of dipolarity, and multipolar Dynamos at larger forcing are associated with the increasing strength of inertia

# **Key question:**

Is the Magnetostrophic balance achievable with todays computer?

#### Roberts number q = Pm / Pr

Carsten Kutzner, Uli Christensen 2003



• E<sub>ma</sub> / \* No dynamo



Morin & Dormy (2005, 2009)





Morin & Dormy (2005, 2009)



Morin & Dormy (2005, 2009)













 $V_{\phi}$ , E = 3 10<sup>-4</sup>, Pm = 18, Ra/Rac = 1.72

$$\begin{cases} E_{\eta} \left( \frac{\partial \mathbf{u}}{\partial t} + (\mathbf{u} \cdot \nabla) \mathbf{u} \right) = -\nabla \pi + E \Delta \mathbf{u} - \mathbf{e}_{z} \wedge \mathbf{u} \\ - Ra T \mathbf{g} + (\nabla \wedge \mathbf{B}) \wedge \mathbf{B}, \\ \nabla \cdot \mathbf{u} = 0, \\ \frac{\partial T}{\partial t} + \mathbf{u} \cdot \nabla T = q \Delta T, \\ \frac{\partial \mathbf{B}}{\partial t} = \nabla \wedge (\mathbf{u} \wedge \mathbf{B}) + \Delta \mathbf{B}, \\ \nabla \cdot \mathbf{B} = 0. \ge 10^{-6} \ge 1 \\ E/Pm = E_{\eta} = \frac{\eta}{2\Omega \mathcal{L}^{2}}, \quad E = \frac{\nu}{2\Omega \mathcal{L}^{2}}, \quad q = \frac{\kappa}{\eta}. \\ \simeq 10^{-7} \simeq 10^{-15} \simeq 10^{-7} \end{cases}$$

 $-2 \frac{\partial \mathbf{u}}{\partial z} \cdot \mathbf{e}_r$ 



 $(\mathbf{\nabla} \times ((\mathbf{\nabla} \times \mathbf{B}) \times \mathbf{B})) \cdot \mathbf{e}_r$ 



(Roberts, GAFD, 1988)





# 3D bifurcation diagram



E=3 10-4



#### Relevant parameter space

Dipolar Viscosity~Coriolis

E Pm

Applied Math: "Weak" Numericists: "Strong"

#### Multipolar / Dynamo wave Inertia~Coriolis

Applied Math: irrelevant to the geodynamo Numericists: "weak"

Dipolar Magnetostrophic Applied Math: "Strong" Numericists: "Stronger" ???



# Toward a distinguished limit

Roberts number q = Pm / Pr



E<sub>ma</sub>
 \* No dynamo

Carsten Kutzner, 2003

# Toward a distinguished limit



# Toward a distinguished limit

 $\mathrm{Pm}_c \sim E^{2/3}$ 

 $\mathrm{Pm}^3 \sim E^2$ 

 $\mathrm{Pm} \sim \varepsilon^2 \qquad E \sim \varepsilon^3$ 

# The Elsasser number



## The Elsasser number



 $\Lambda = \frac{B^2}{2\Omega\rho_0\mu_0\eta}$ 

# The Elsasser number

Ra Ra/Ra<sub>c</sub> Pm Rm  $\Lambda \Lambda/(Rm \ell_B)$ 

| 105 | 1.7 | 18 | 195 | 1.14 | 0.06 |
|-----|-----|----|-----|------|------|
| 105 | 1.7 | 18 | 145 | 13.2 | 1.11 |
| 125 | 2.0 | 18 | 207 | 24.0 | 1.54 |
| 150 | 2.5 | 12 | 165 | 23.7 | 1.65 |

 $\Lambda = \frac{2 E_m}{V_c} Pm E$ 

# CLAIMS

None of the published spherical dynamo models correspond to a force balance relevant to the geodynamo!
This force balance can be

approached in numerical models.

# **OPENED ISSUES**

- All terms are of similar amplitude, Yet well identified balances seem to emerge. Why? - How do these models evolve at lower E, and lower Pm? - Is it possible to sustain dynamo action for Ra<Ra?
## Glatzmaier & Roberts, 1995

